

Coupled atmosphere-ocean data assimilation in the presence of model error

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Introduction

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Model error in the coupled system **restricts the window length** which can be used with 4D-Var to something shorter than the optimal window length in an uncoupled Ocean DA scheme.

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Aim of work is to understand how **different coupling strategies** react as window length is extended and the model error becomes more significant.

Coupling strategies (recap)

Each method is based on incremental 4D-Var. In each case the first guess comes from a coupled forecast.

Strongly coupled DA: uses the coupled model in both the inner and outer loops.

Weakly coupled DA: uses the coupled model in the outer loop but the inner loop is uncoupled.

Uncoupled DA: uses the uncoupled models in both the outer and inner loops. BCs at the interface are prescribed externally.

Model error in 4D-Var

- In 4D-Var have the assumption that the background state, \mathbf{x}^b , and the observations, $\hat{\mathbf{y}}$, are consistent with

$$\mathbf{x}^b \sim N(\mathbf{x}_0^t, \mathbf{B}) \quad \text{and} \quad \hat{\mathbf{y}} \sim N(\hat{\mathbf{y}}^t, \hat{\mathbf{R}})$$
$$\text{and} \quad \hat{\mathbf{y}}^t = \hat{\mathcal{H}}(\mathbf{x}_0^t)$$

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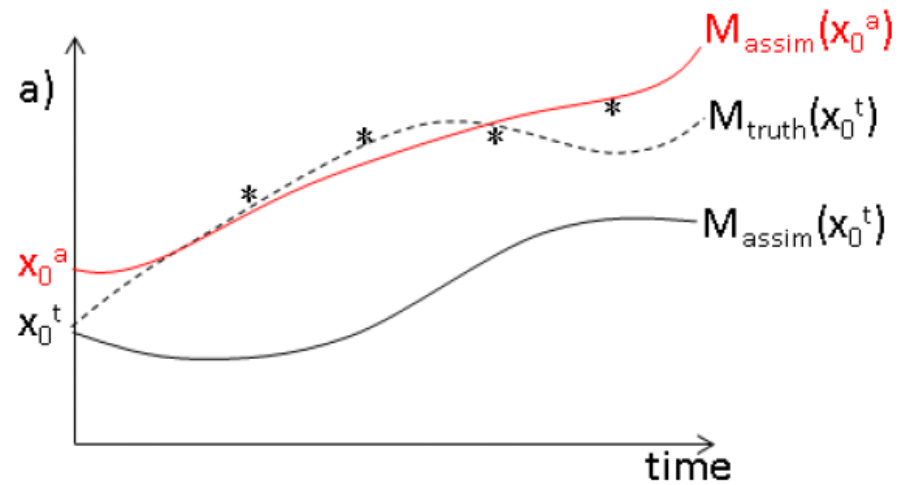
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- However if model error becomes significant then this last assumption breaks down, and instead:

$$\begin{aligned} \hat{\mathbf{y}}^t &= \hat{\mathcal{H}}^t(\mathbf{x}_0^t) \\ &= \hat{\mathcal{H}}(\mathbf{x}_0^t) + \epsilon^{\hat{\mathcal{H}}} \quad \text{where} \quad \epsilon^{\hat{\mathcal{H}}} \in \mathbb{R}^{\hat{p} \times 1} \end{aligned}$$

Model error in 4D-Var

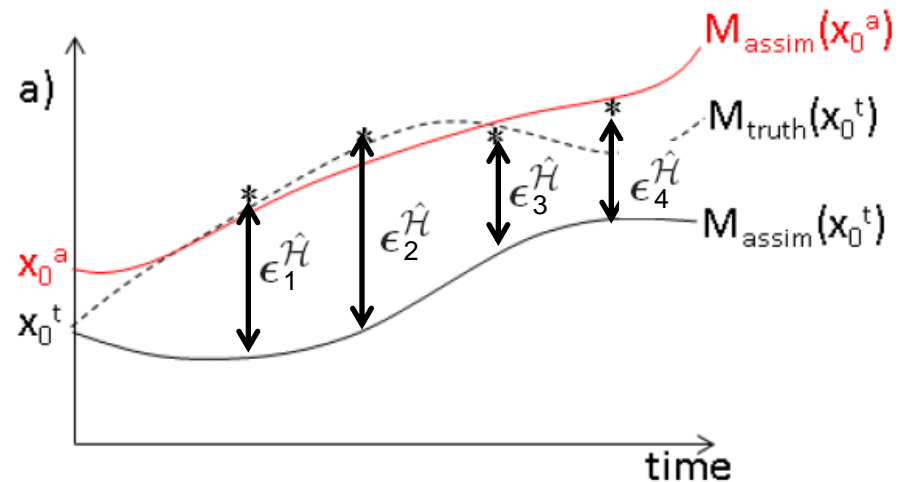


Model error in 4D-Var

If the model error is unaccounted for it has the effect of increasing the analysis error covariances

$$E[\epsilon^a(\epsilon^a)^T] = \mathbf{P}_{nm}^a + \mathbf{K}E[\epsilon^{\hat{H}}(\epsilon^{\hat{H}})^T]\mathbf{K}^T$$

where \mathbf{P}_{nm}^a is the analysis error covariance matrix when no model error is present.

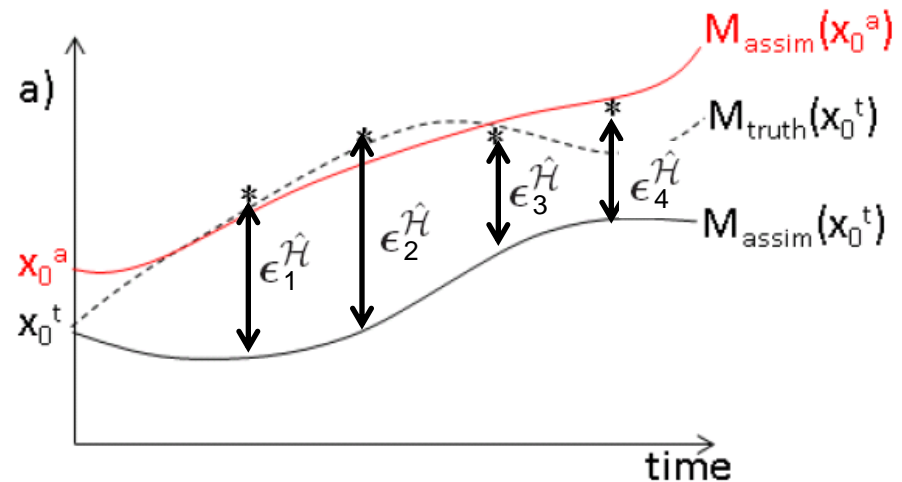


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And if the model error is biased then the analysis error will also be biased

$$E[\epsilon^a] = \mathbf{K}E[\epsilon^{\hat{\mathcal{H}}}]$$

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Strongly coupled $\epsilon^{\hat{\mathbf{H}}} = \hat{\mathcal{H}}^t(\mathbf{x}_0^t) - \hat{\mathcal{H}}^c(\mathbf{x}_0^g) - \hat{\mathbf{H}}^c|_{\mathbf{x}^g}(\mathbf{x}_0^t - \mathbf{x}_0^g)$

Weakly coupled $\epsilon^{\hat{\mathbf{H}}} = \hat{\mathcal{H}}^t(\mathbf{x}_0^t) - \hat{\mathcal{H}}^c(\mathbf{x}_0^g) - \hat{\mathbf{H}}^{\text{uc}}|_{\mathbf{x}^g}(\mathbf{x}_0^t - \mathbf{x}_0^g)$

uncoupled $\epsilon^{\hat{\mathbf{H}}} = \hat{\mathcal{H}}^t(\mathbf{x}_0^t) - \hat{\mathcal{H}}^{\text{uc}}(\mathbf{x}_0^g) - \hat{\mathbf{H}}^{\text{uc}}|_{\mathbf{x}^g}(\mathbf{x}_0^t - \mathbf{x}_0^g)$

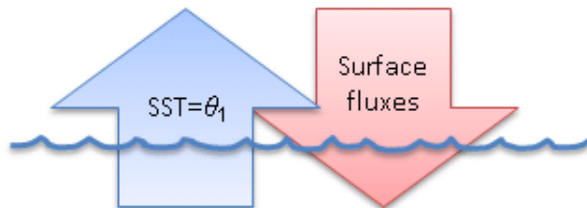
Model set up

- generation of model error

Atmosphere

T_1	q_1	u_1	v_1
T_2	q_2	u_2	v_2
T_3	q_3	u_3	v_3

T_N	q_N	u_N	v_N
-------	-------	-------	-------



θ_1	s_1	u_1	v_1
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θ_{M-2}	s_{M-2}	u_{M-2}	v_{M-2}
θ_{M-1}	s_{M-1}	u_{M-1}	v_{M-1}
θ_M	s_M	u_M	v_M

Ocean

The model error is assumed to be complex and from multiple sources.

Atmosphere

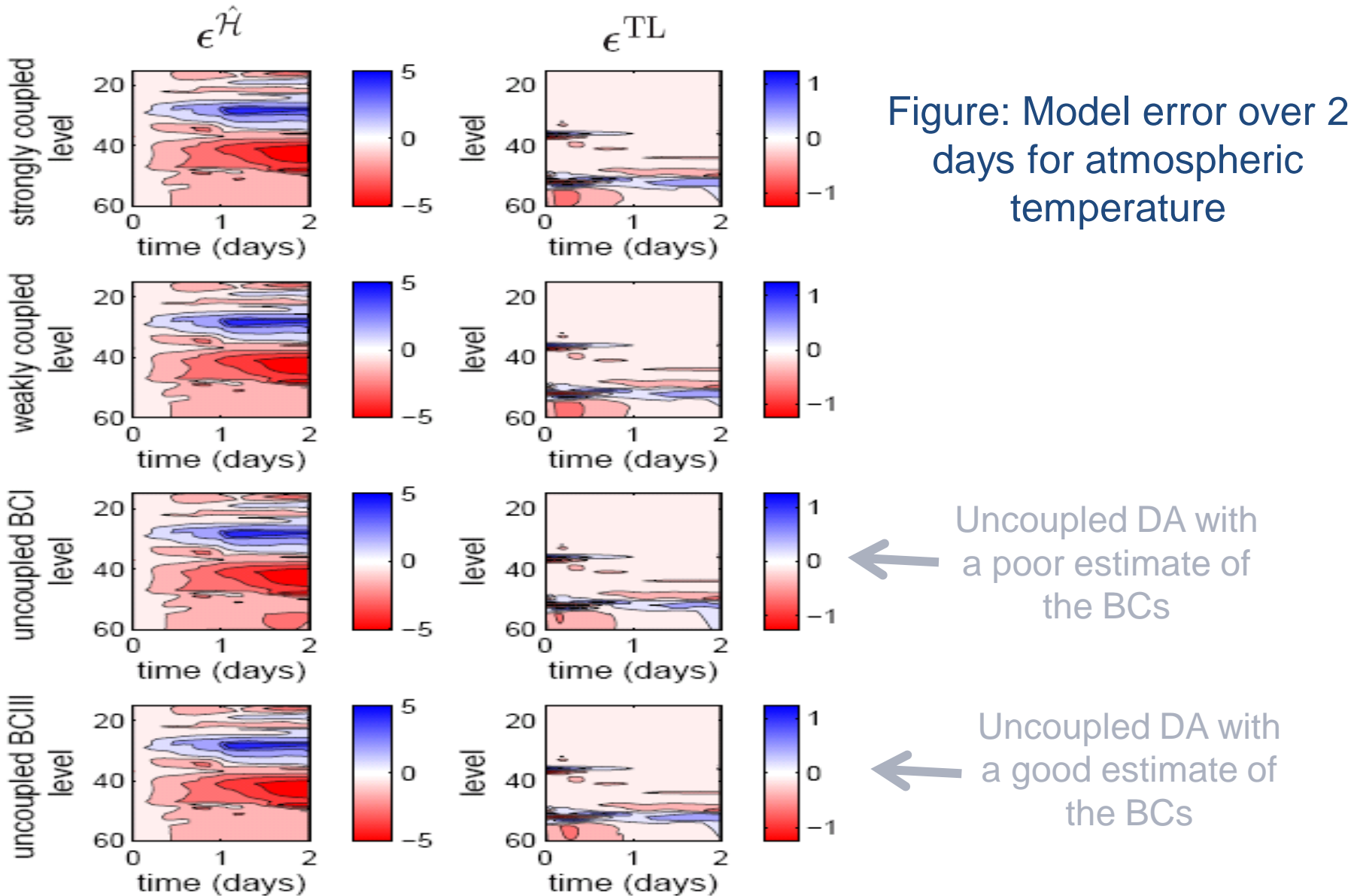
Assimilation model has missing physics and a bias in the large scale forcing.

Ocean

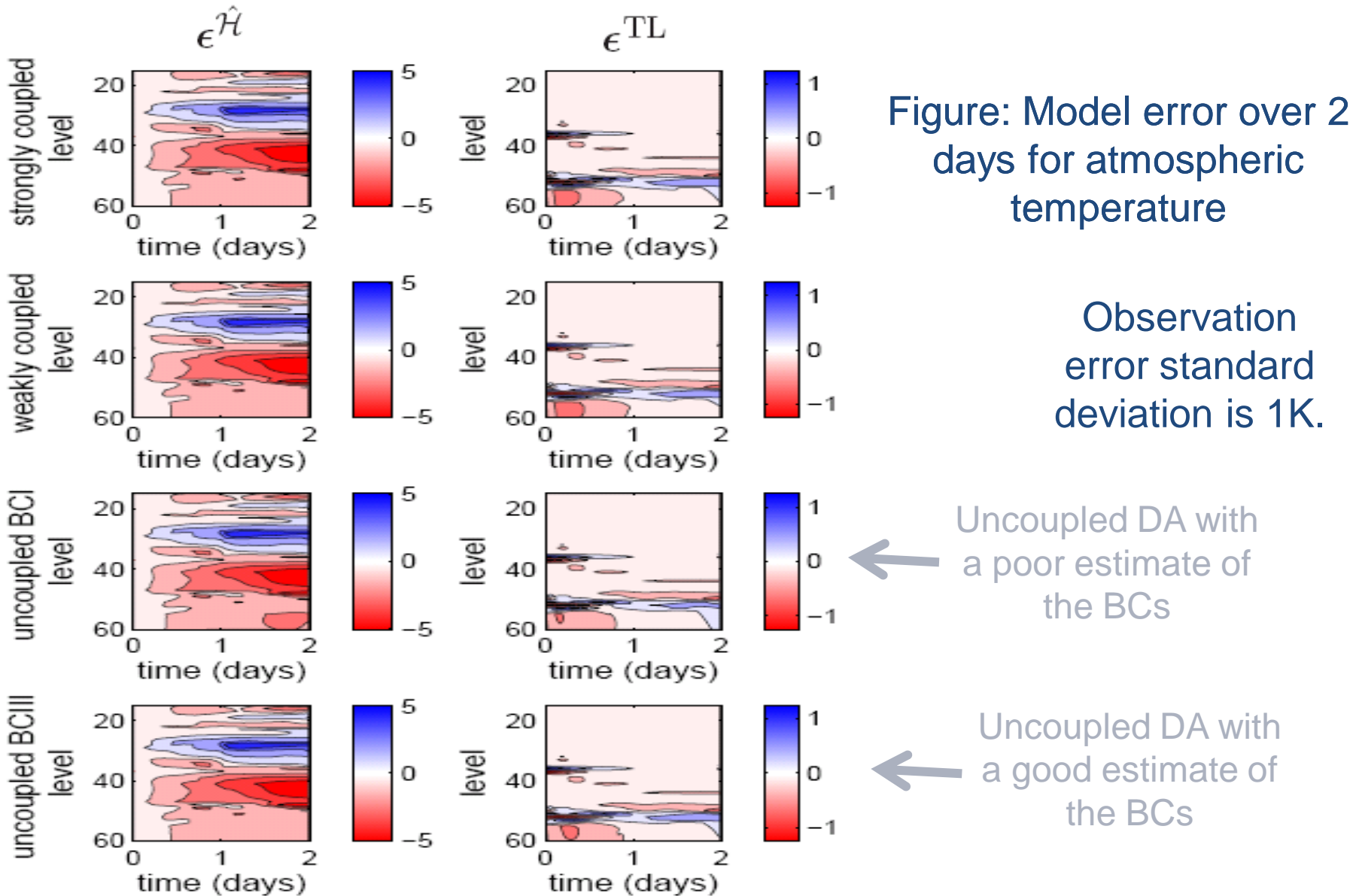
Assimilation model has perturbed diffusion parameters.

Results are shown for the July 2014 case study for a point in the NW Pacific.

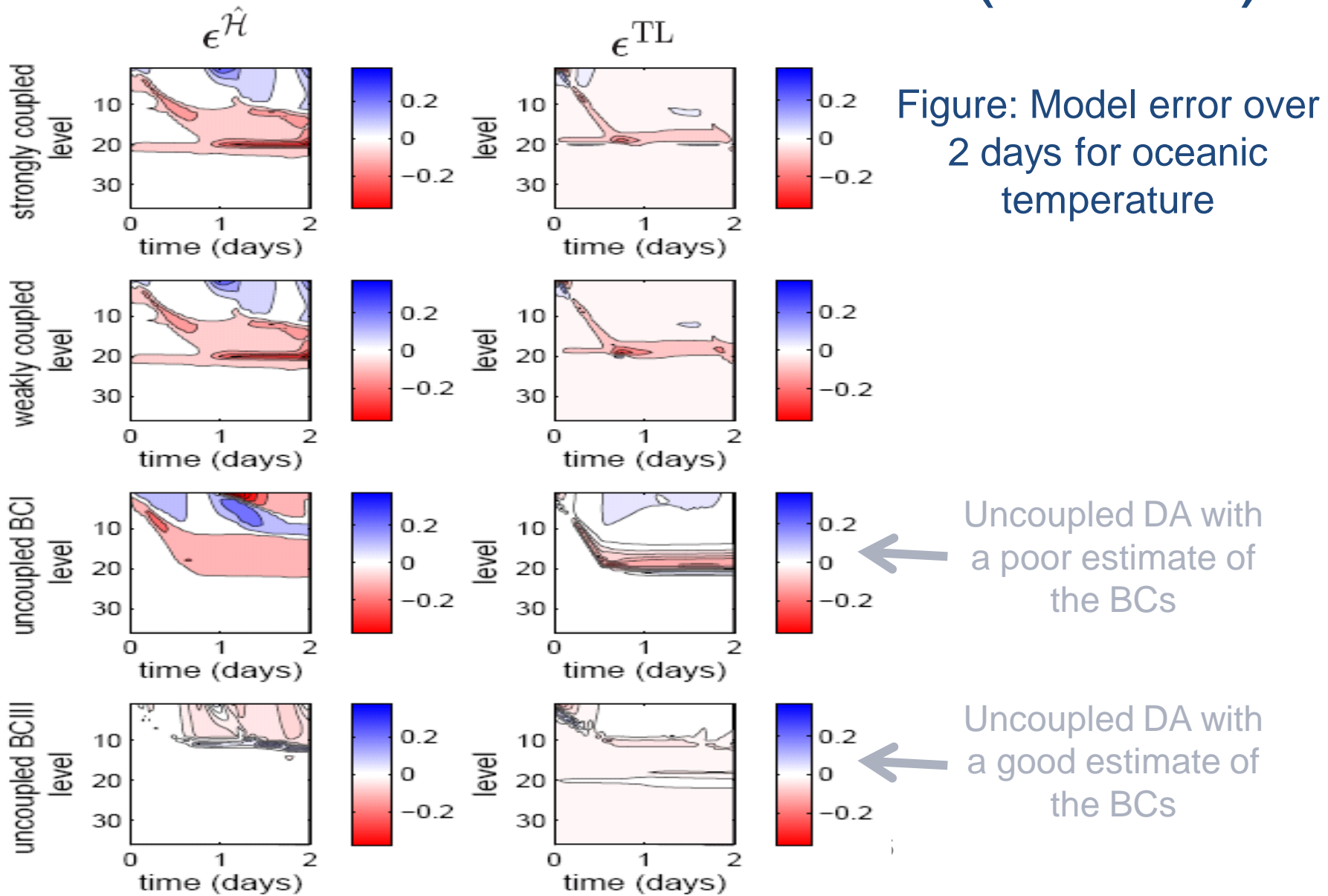
Model error in the atmosphere (no DA)



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Model error in the ocean (no DA)



Model error in the ocean (no DA)

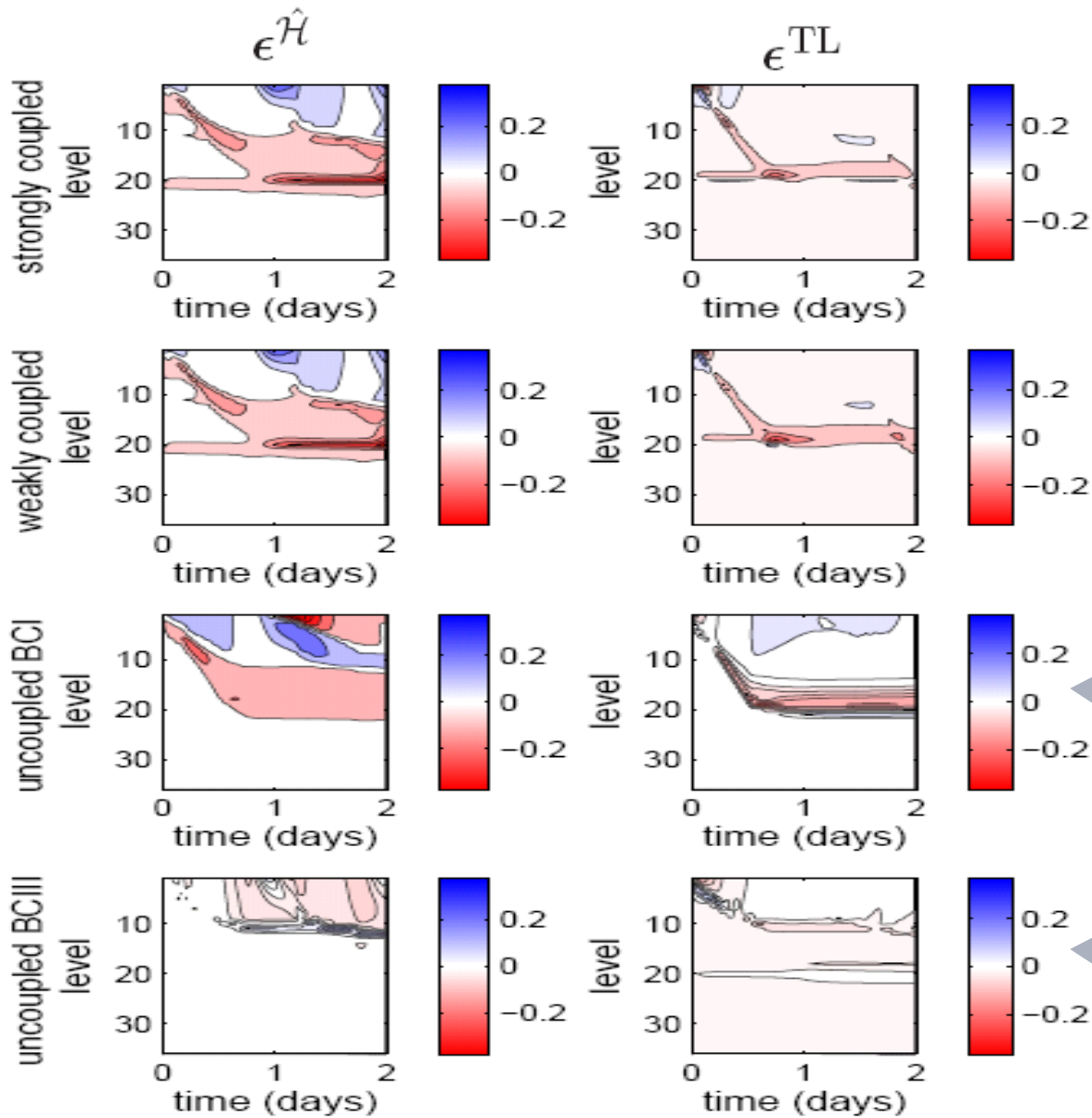


Figure: Model error over 2 days for oceanic temperature

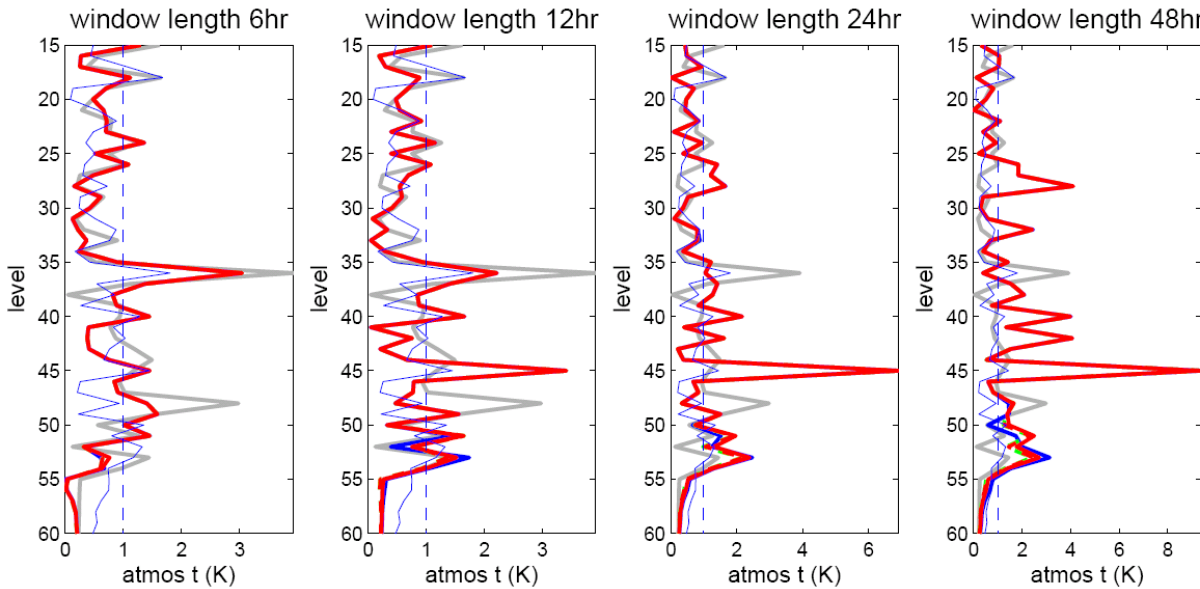
Observation error standard deviation is 0.1K.

← Uncoupled DA with a poor estimate of the BCs

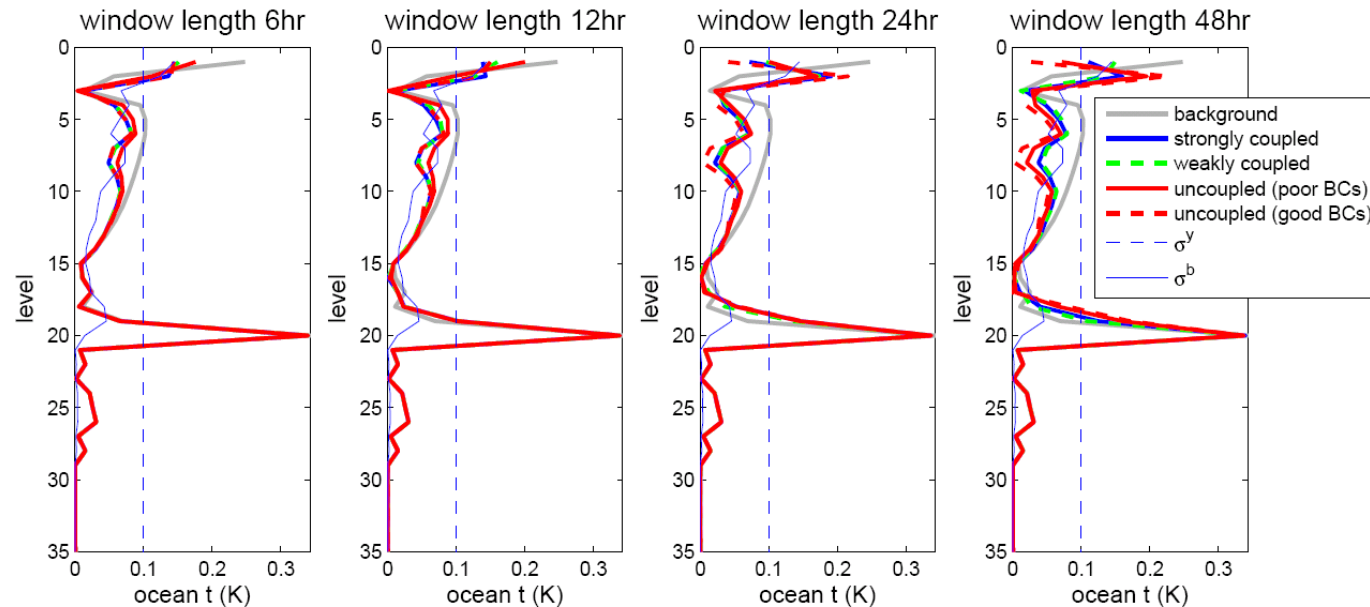
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Absolute analysis error ($|x^a - x^t|$)

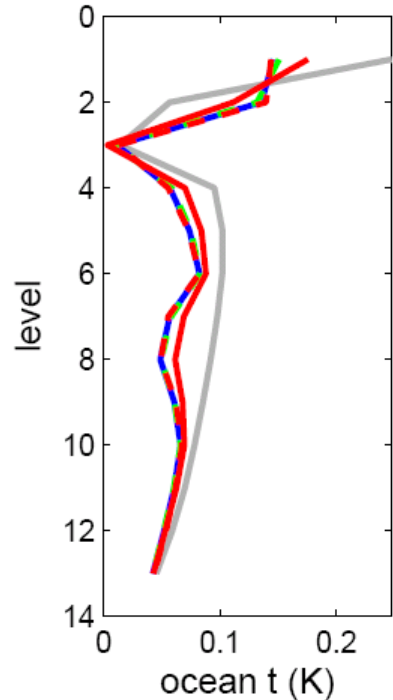
Figures: Absolute error in temperature at the initial time, in the atmosphere (left) and ocean (below).



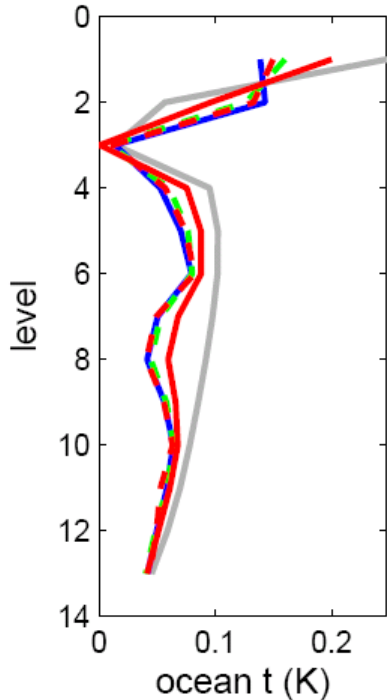
- Atmospheric observations every 3 hours, ocean observations every 6 hours.
- Each level is observed.
- B is diagonal



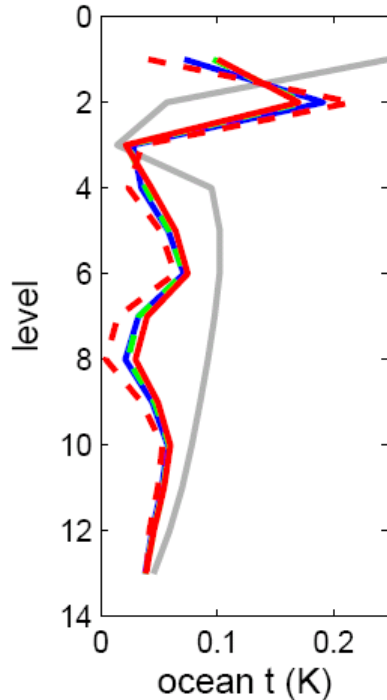
window length 6hr



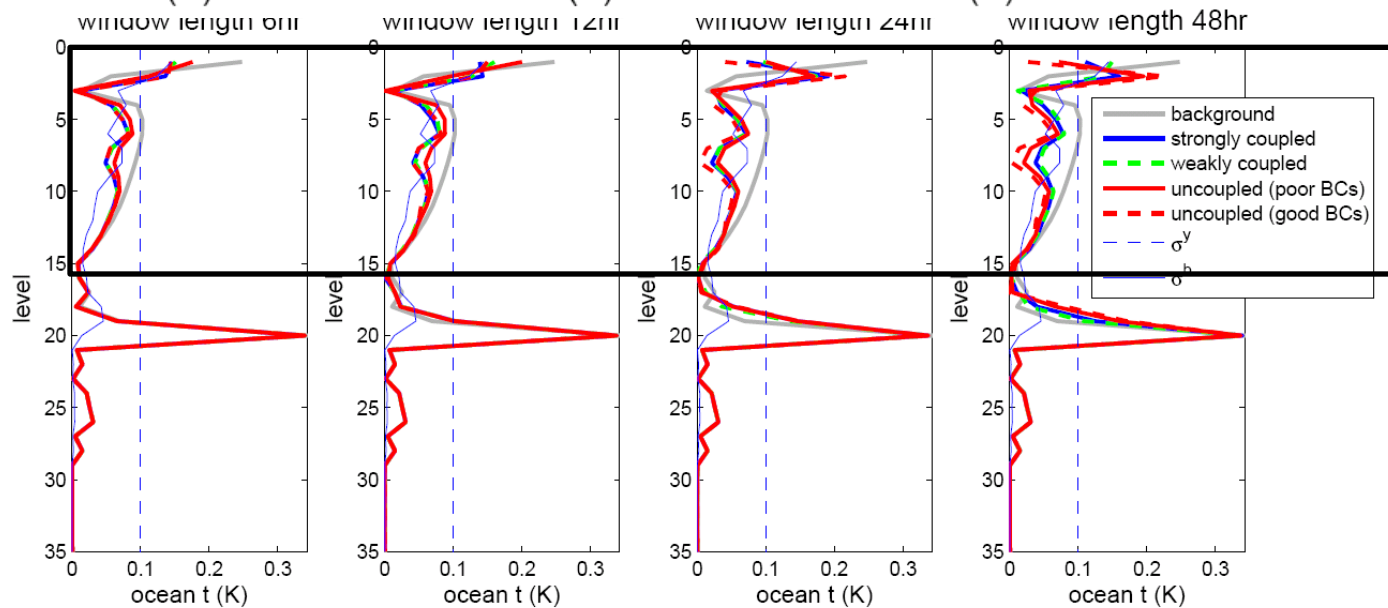
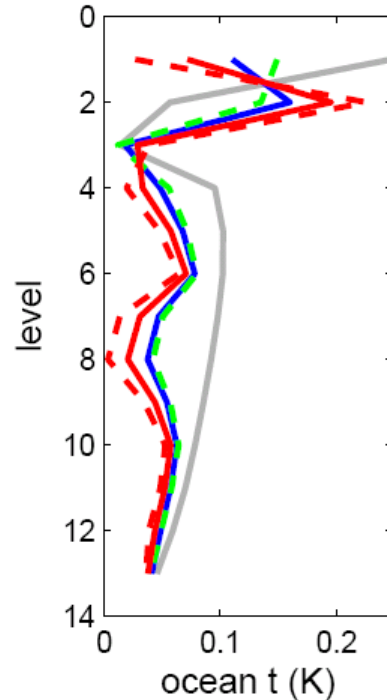
window length 12hr



window length 24hr



window length 48hr



Forecast error in the atmosphere

$$(M^t(x^t) - M(x^a))$$

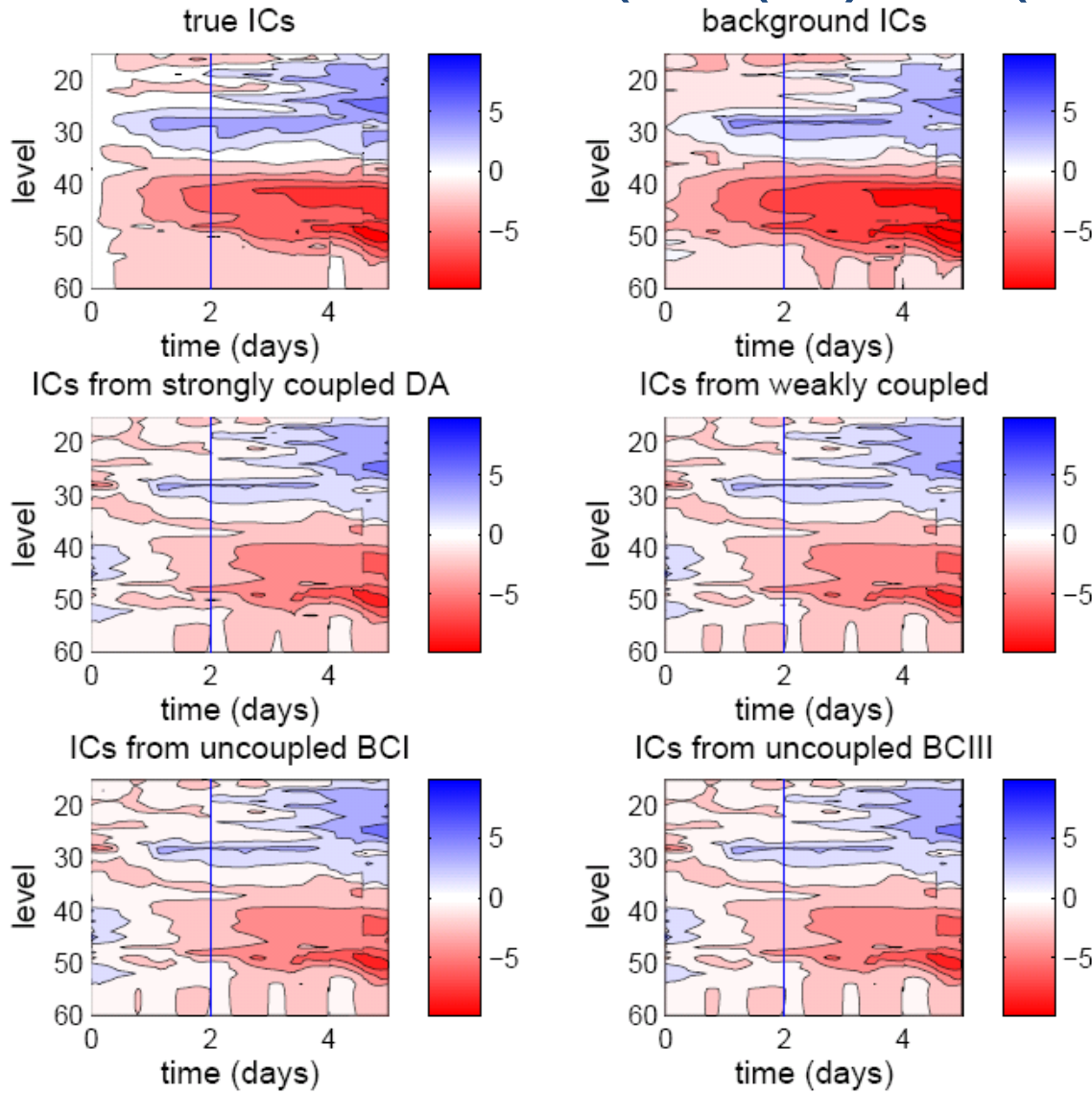


Figure: Forecasts of atmospheric temperature using coupled model initialised using different analyses computed using a 2 day assimilation window

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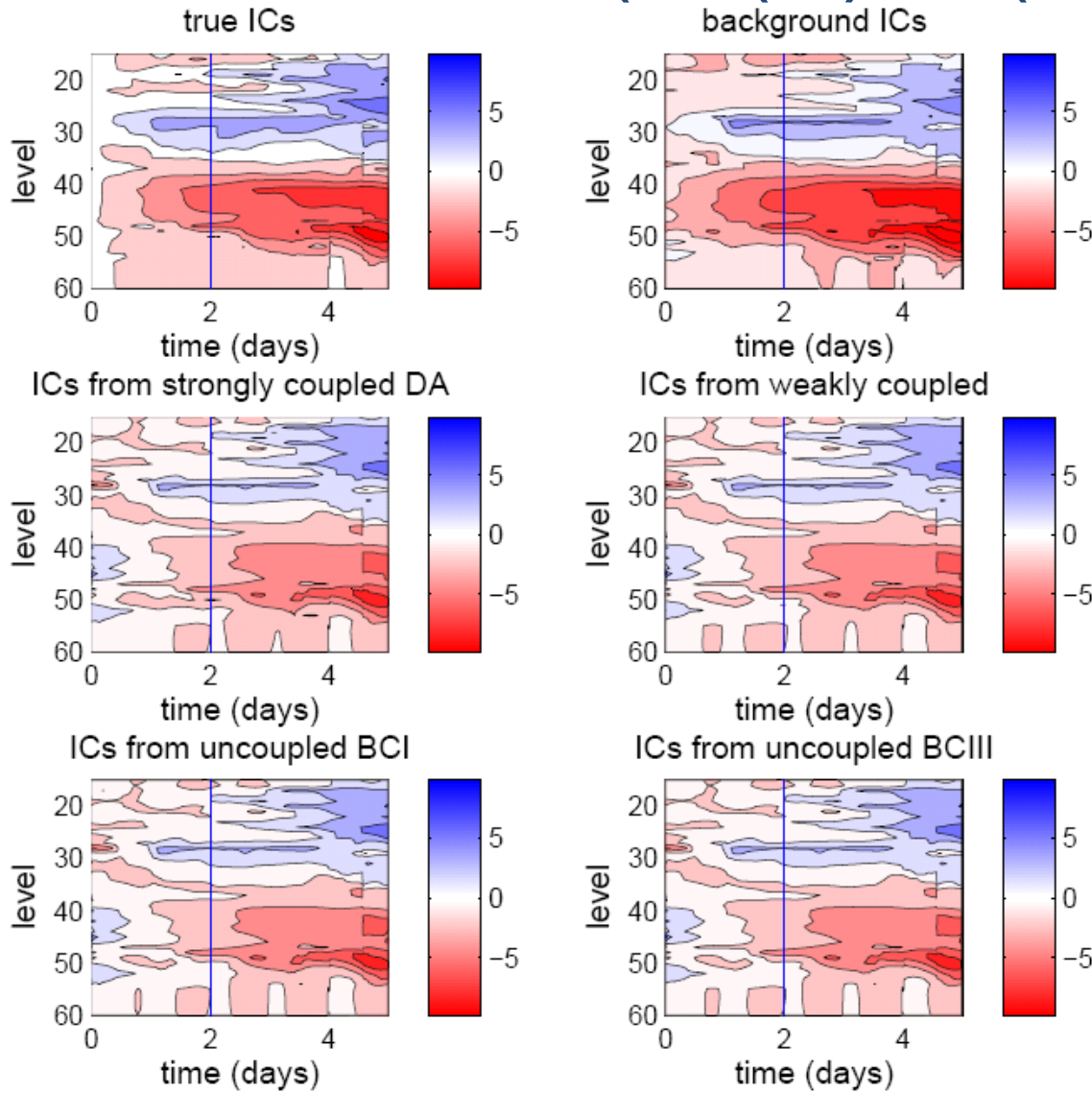
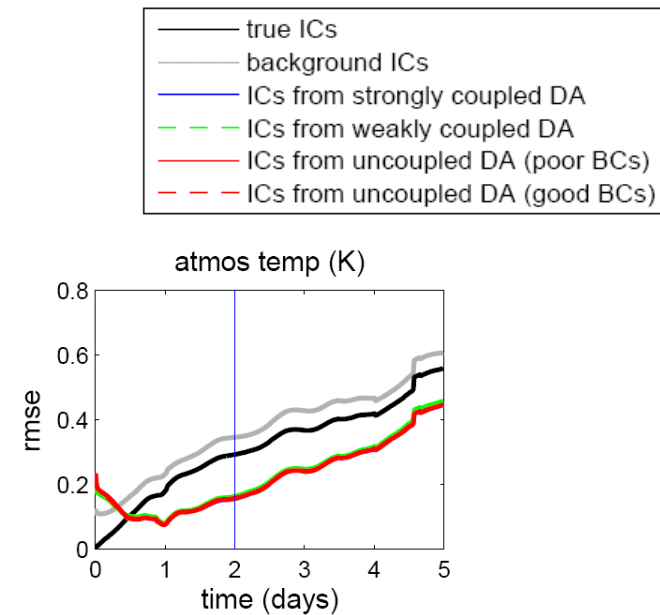


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Forecast error in the ocean

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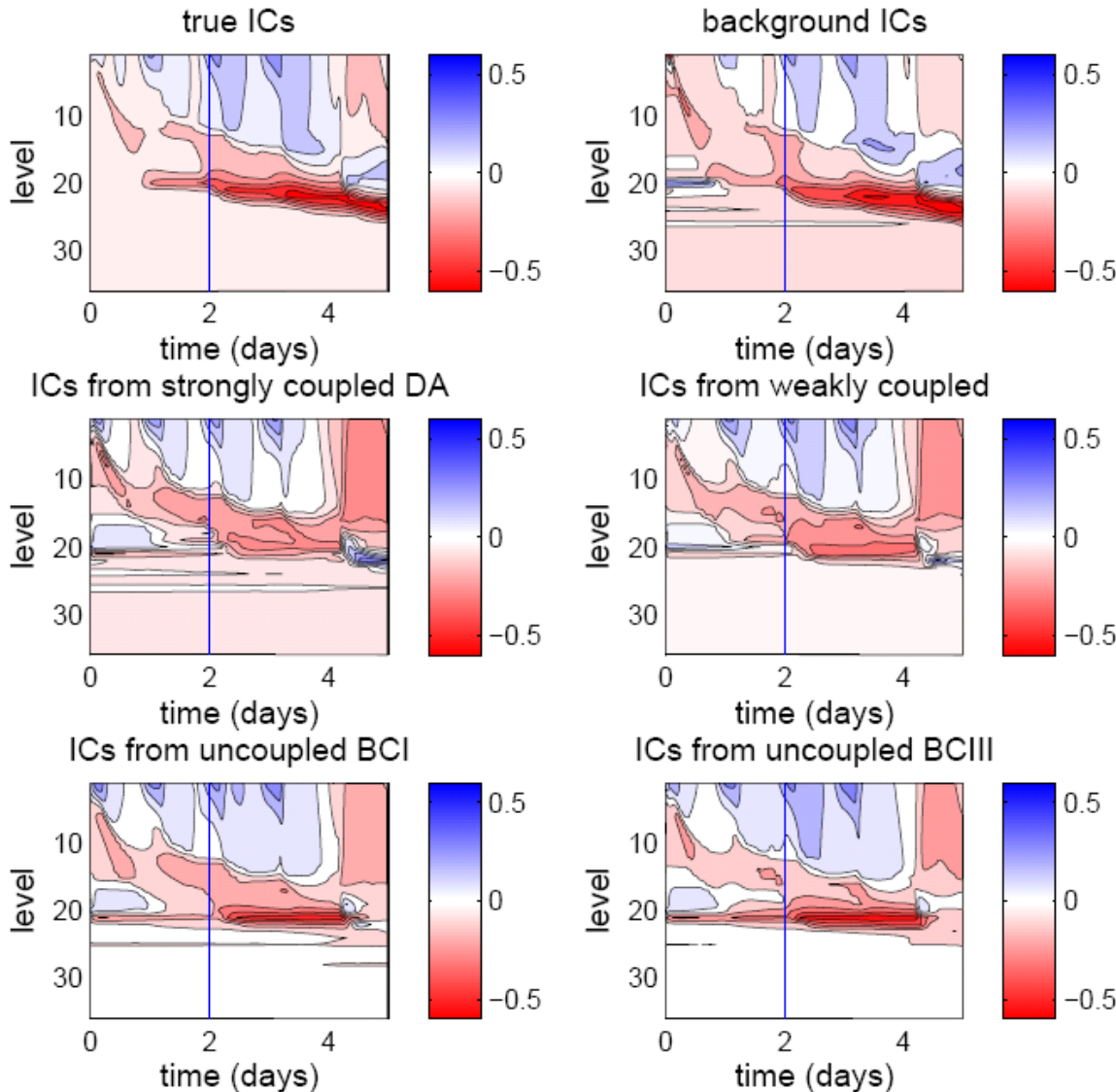


Figure: Forecasts of oceanic temperature using coupled model, initialised using different analyses computed using a 2 day assimilation window

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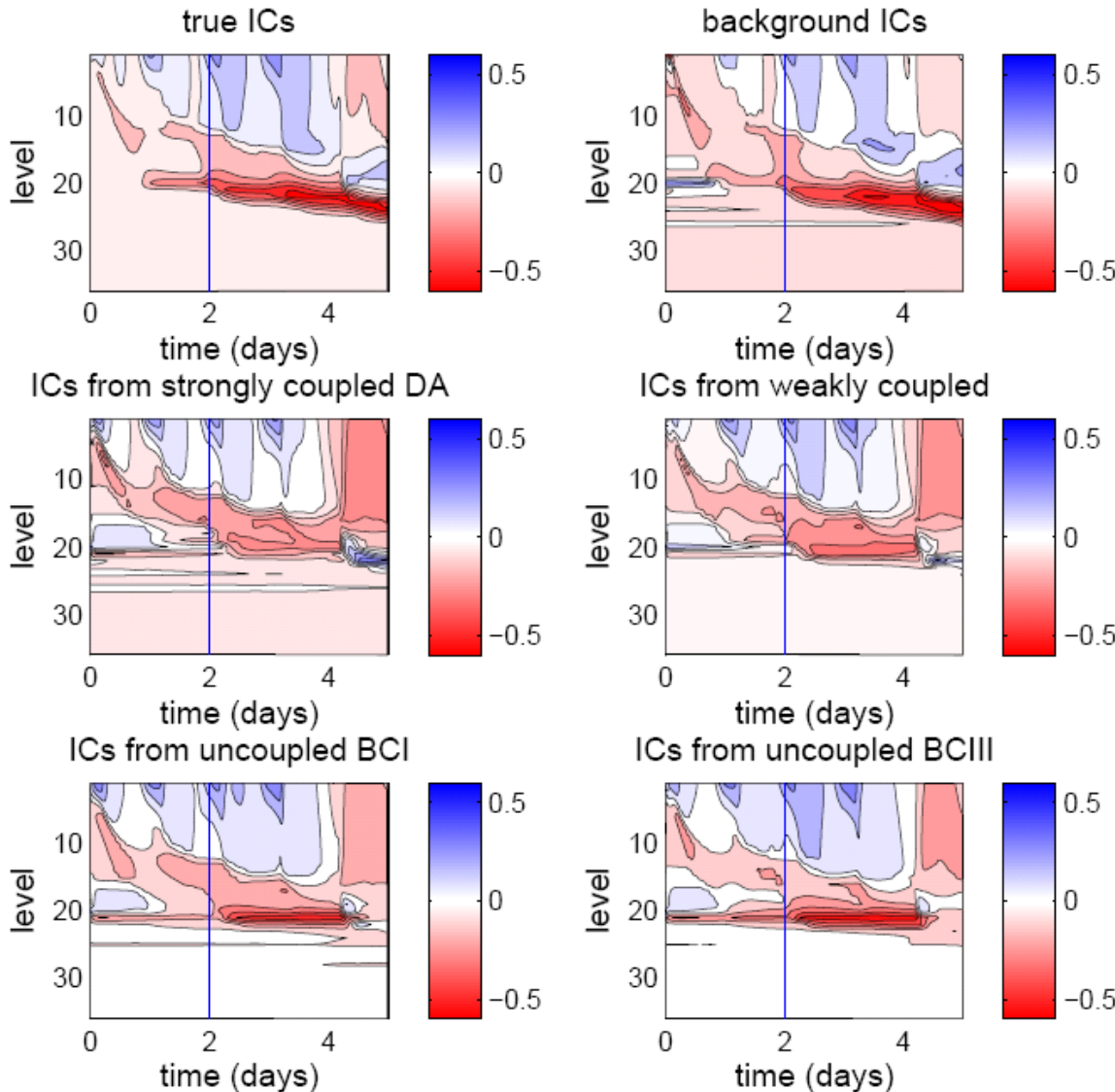
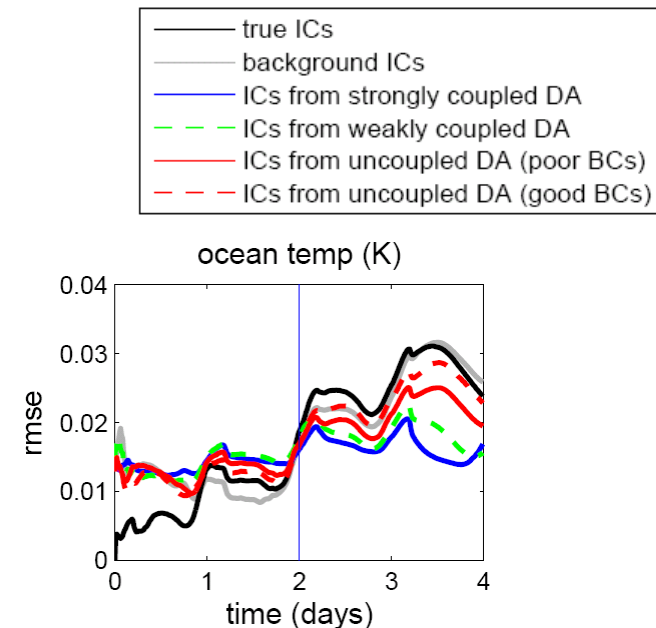


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Summary- 2 day window length

Strongly coupled DA

- poor analysis due to observations being inconsistent with assimilation model
- Produces a more balanced initial state and results in the best forecast beyond a day.

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Strongly coupled DA	Uncoupled DA
<ul style="list-style-type: none">•poor analysis due to observations being inconsistent with assimilation model•Produces a more balanced initial state and results in the best forecast beyond a day.	<ul style="list-style-type: none">•ocean analysis is least affected by model error originating from the surface and therefore is the most accurate.•assimilation model is inconsistent with the forecast model• forecasts initialised with the uncoupled analysis exhibit greatest shock and fastest error growth.

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Strongly coupled DA	Uncoupled DA	Weakly coupled DA
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Conclusions

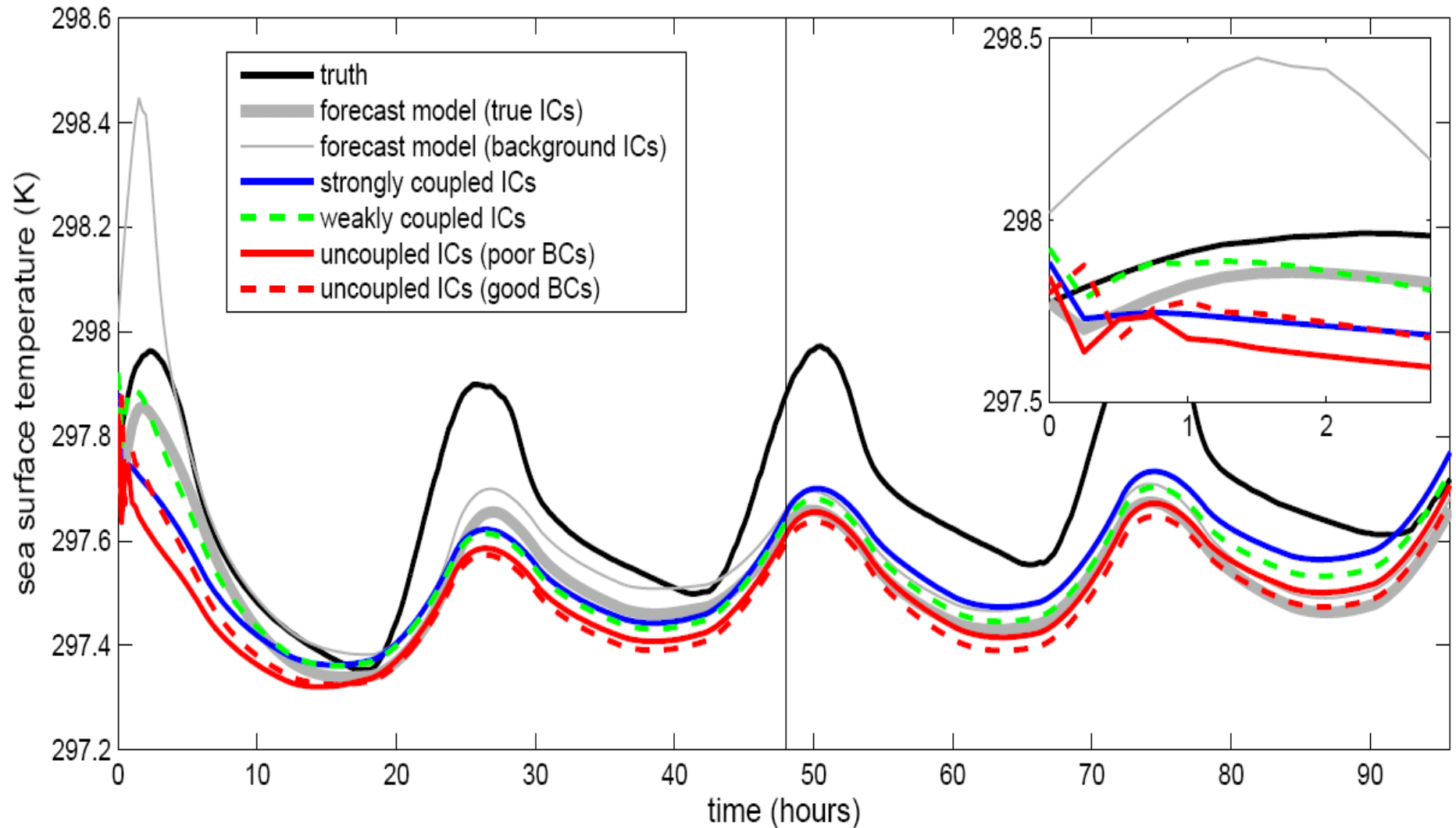
- The coupling of the atmosphere and ocean can amplify the presence of model error.
- The effect of model error on the analysis depends on the coupled DA scheme used.
- Strongly coupled DA has been shown to be able to provide an analysis consistent with the forecast model at the expense of the accuracy of the ocean analysis.
- To improve the utility of strongly coupled DA need to be able to account for model error in the assimilation to allow for the window length to be extended.
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Thank you for listening

Initialisation shock



Initialisation shock- reduced observations

