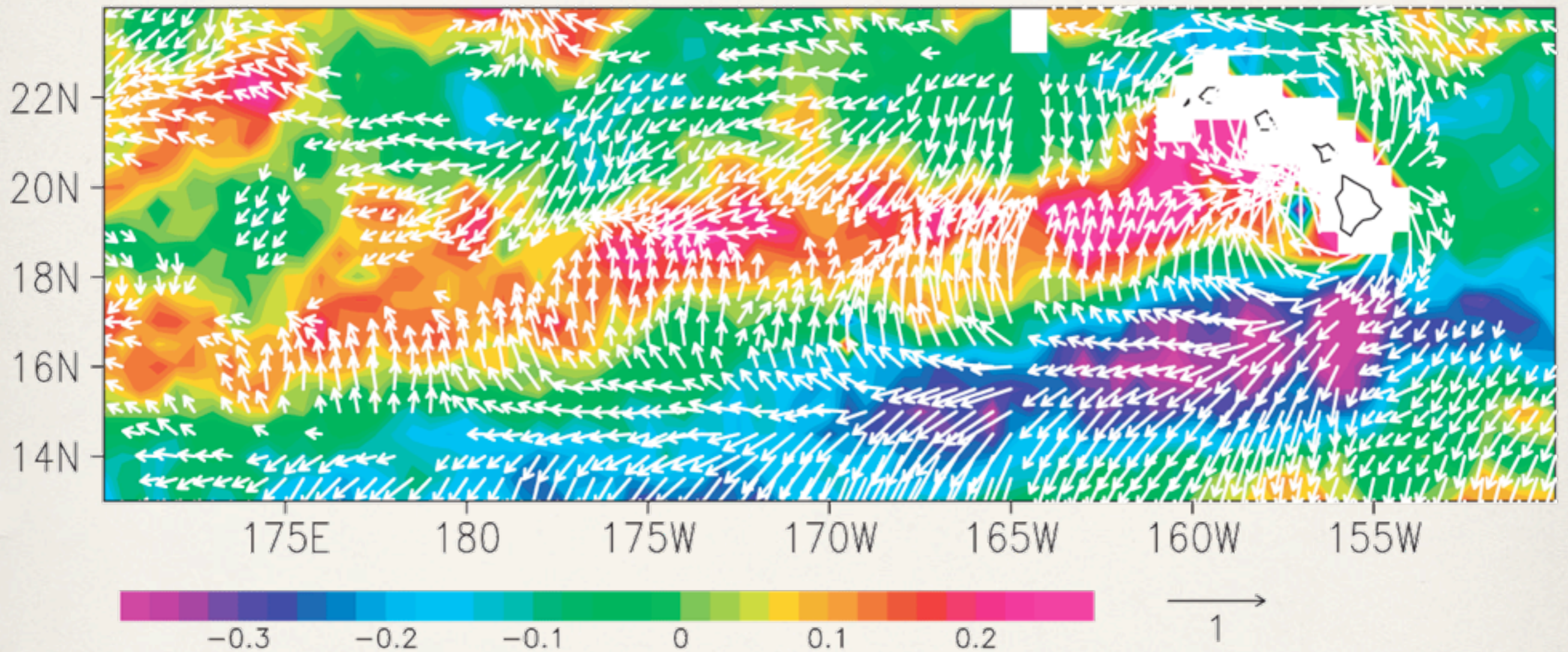


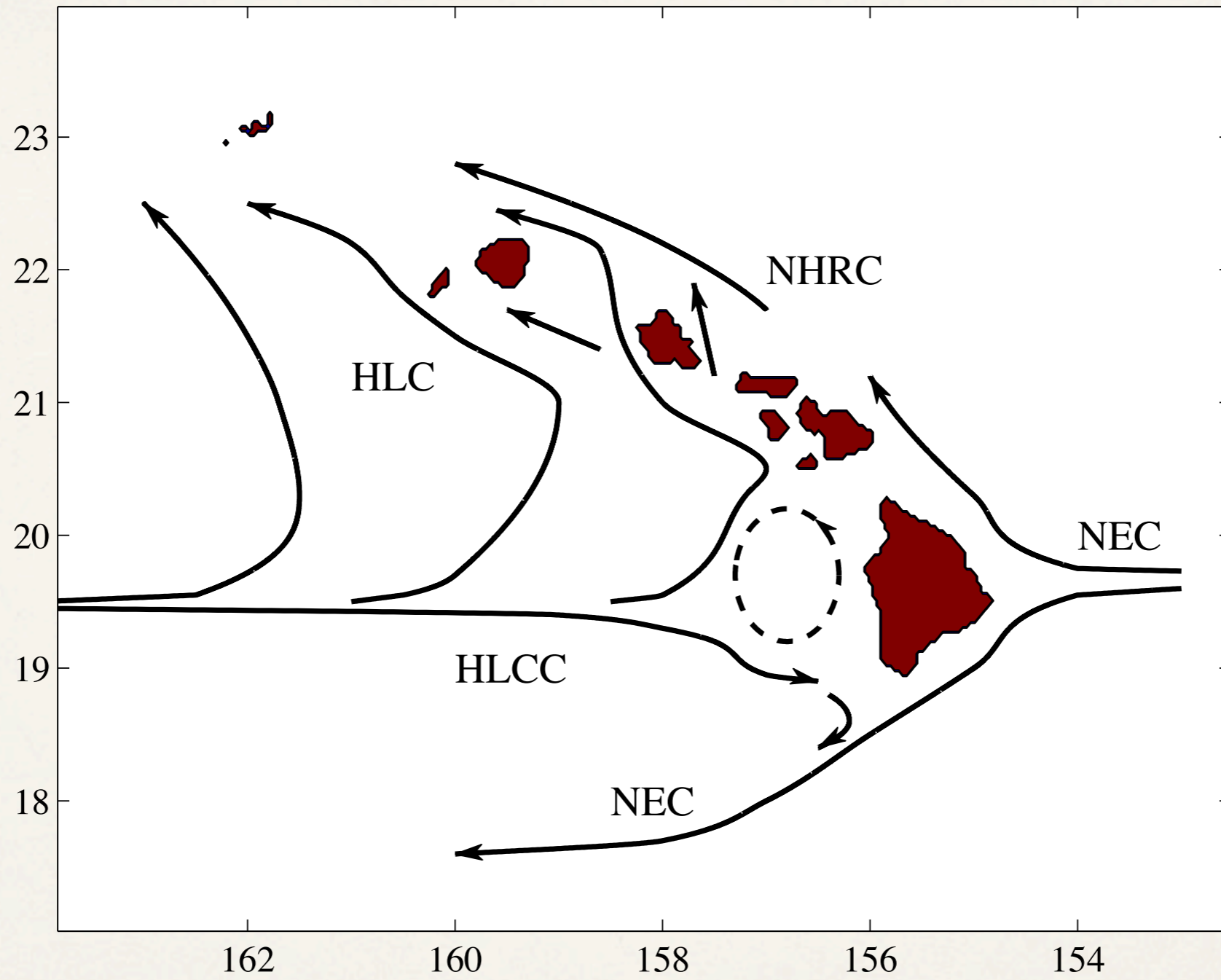


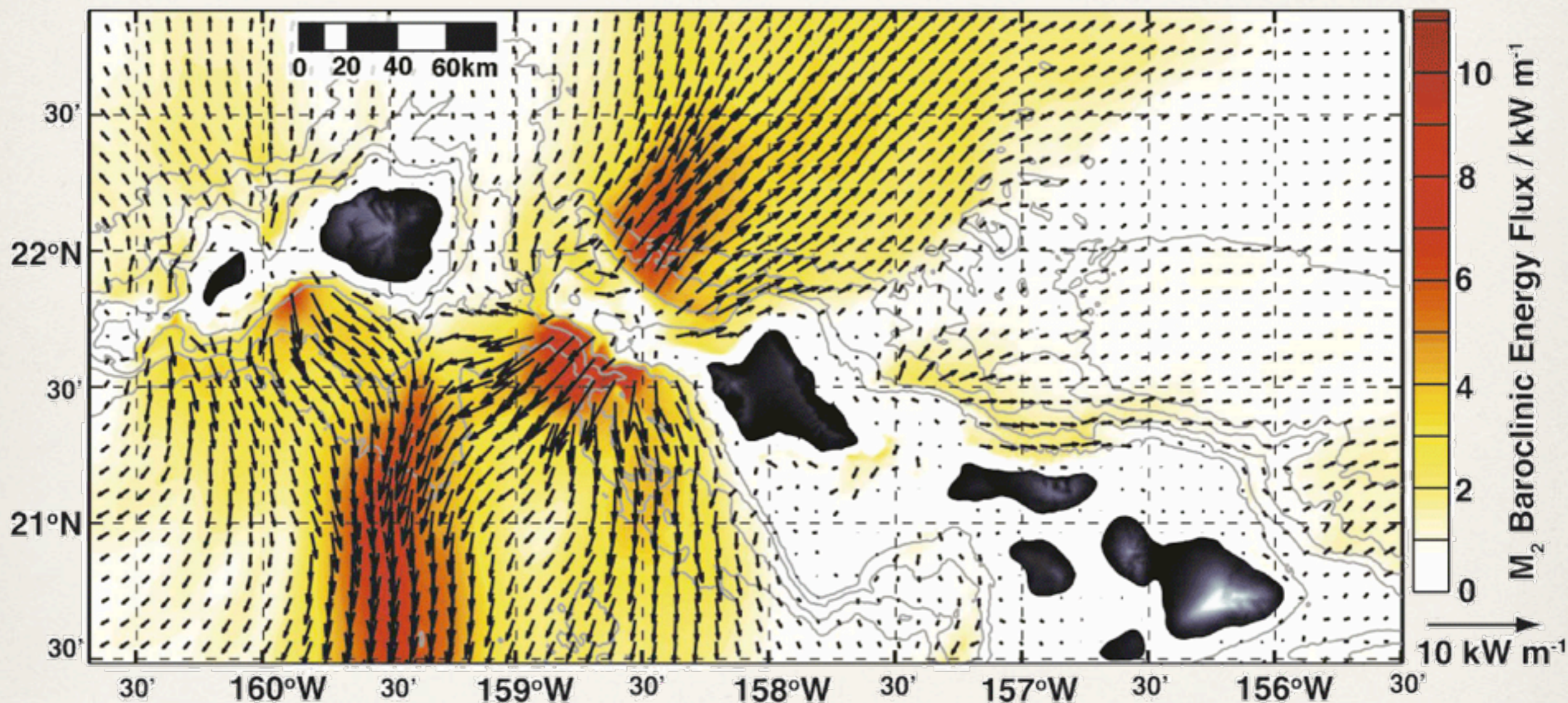
Quantifying the role of observations in ocean state estimation.

Brian Powell, University of Hawaii



S.-P. Xie, W. Liu, Q. Liu, and M. Nonaka. Far-Reaching Effects of the Hawaiian Islands on the Pacific Ocean-Atmosphere System. *Science*, 292(5524):2057–2060, 2001.



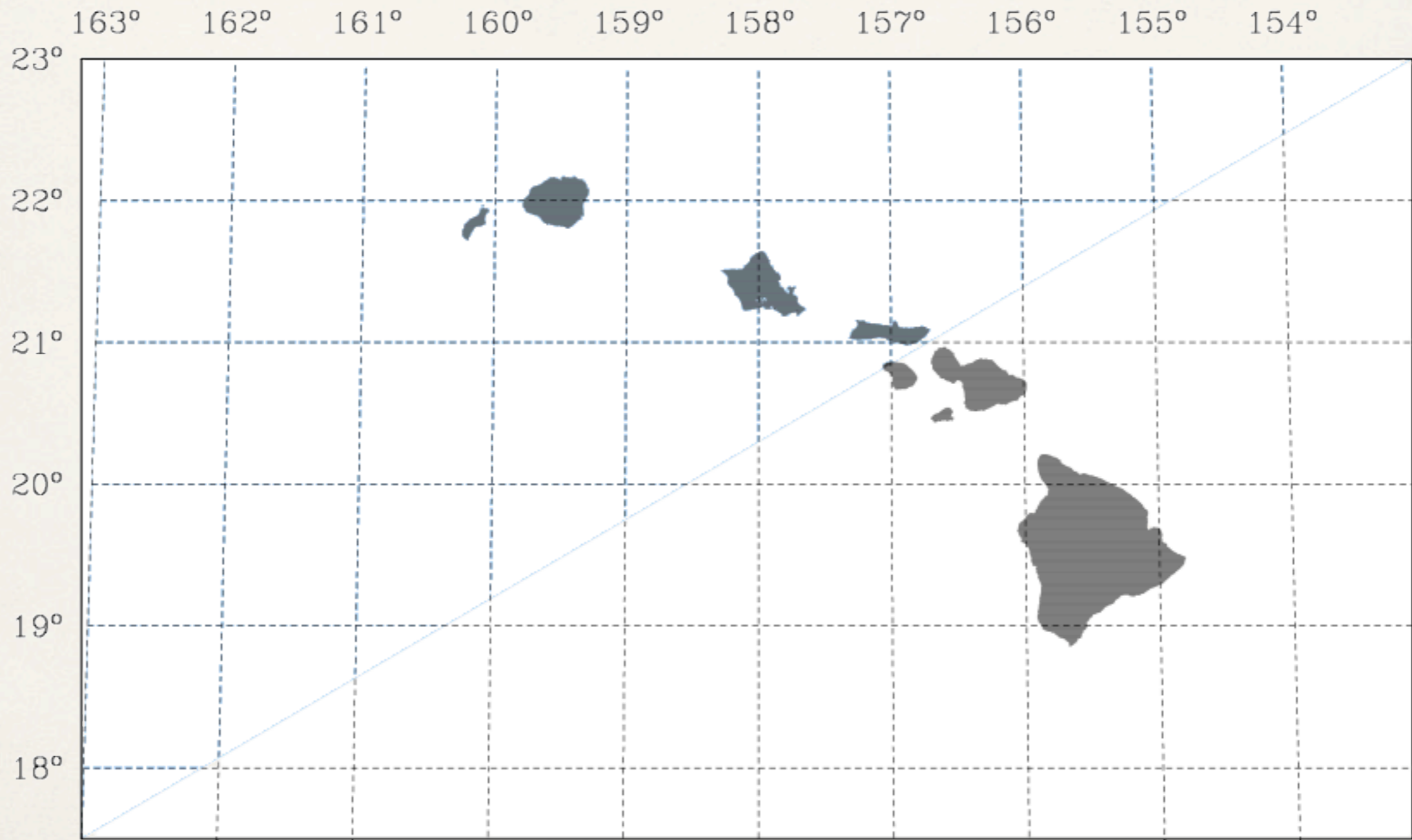


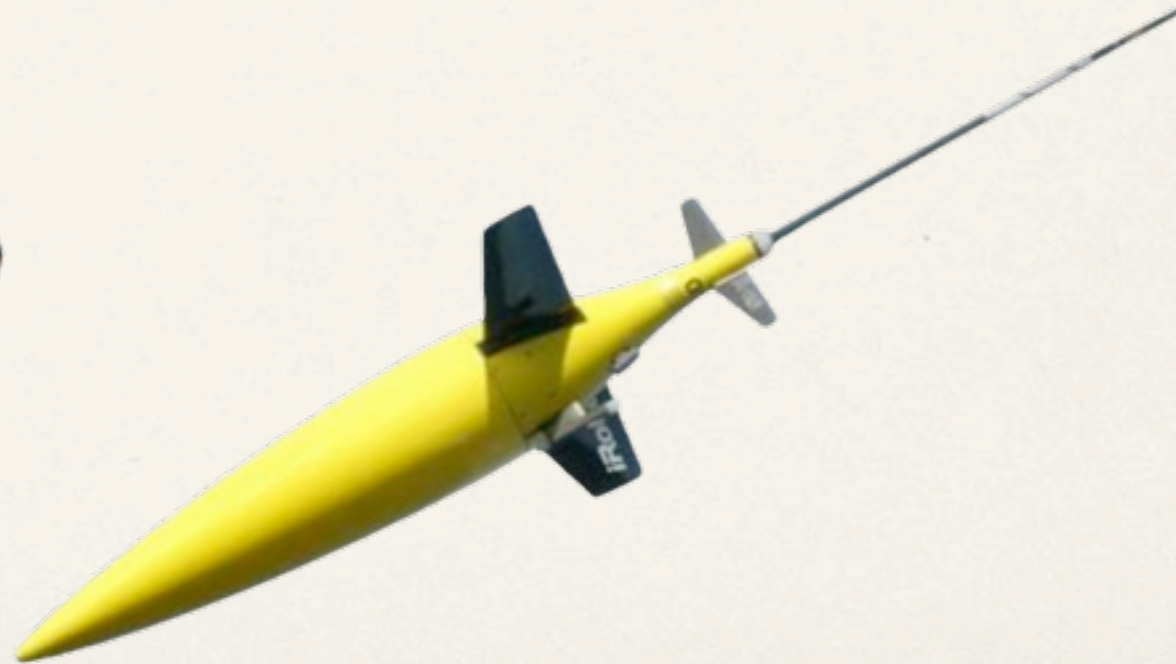
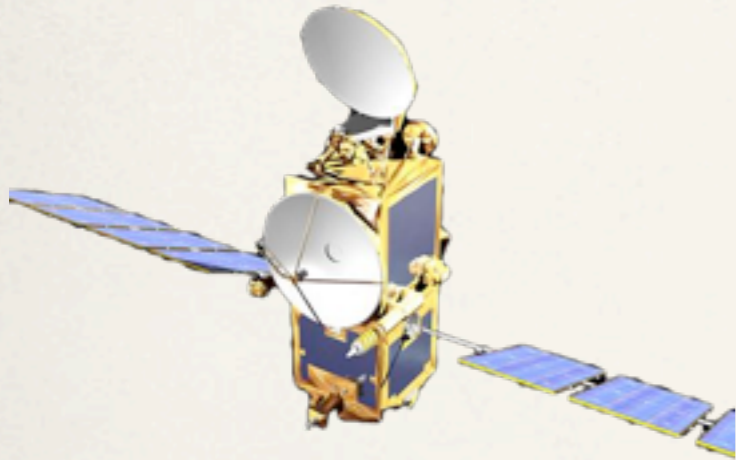
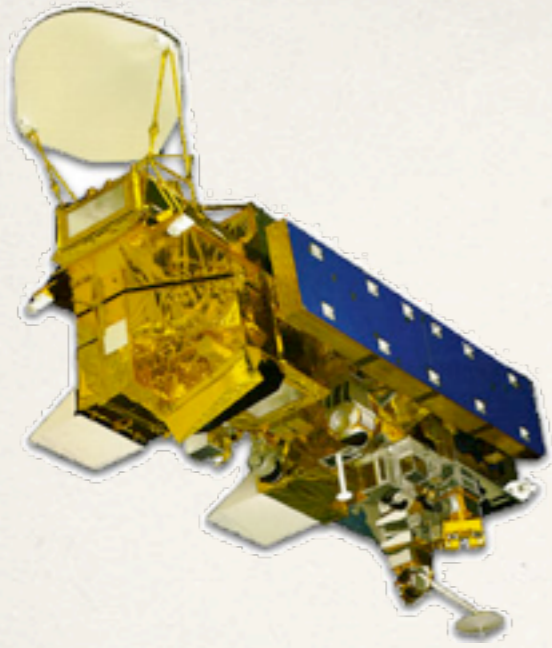
G. S. Carter, M. A. Merrifield, J. M. Becker, K. Katsumata, M. C. Gregg, D. S. Luther, M. D. Levine, T. J. Boyd, and Y. L. Firing. Energetics of M₂ Barotropic-to-Baroclinic Tidal Conversion at the Hawaiian Islands. *J. Phys. Oceanogr.*, 38:2,205–2,223, 2008.



PacOOS Pacific Islands Ocean Observing System

IN THE SCHOOL OF OCEAN AND EARTH SCIENCE AND TECHNOLOGY AT THE UNIVERSITY OF HAWAI'I AT MĀNOA







Observation	Count	Percent
HOT Temperature	4,982	0.02%
HOT Salt	4,982	0.02%
Argo Temp	15,212	0.06%
Argo Salt	15,212	0.06%
Seaglider Temperature	220,266	0.83%
Seaglider Salt	220,266	0.83%
SST	25,201,519	94.50%
SSH	985,731	3.70%
Total	26,668,170	



❖ Recall, $\mathbf{x}_a = \mathbf{x}_b + \mathbf{K} (\mathbf{y} - \mathbf{H}\mathbf{x}_b)$

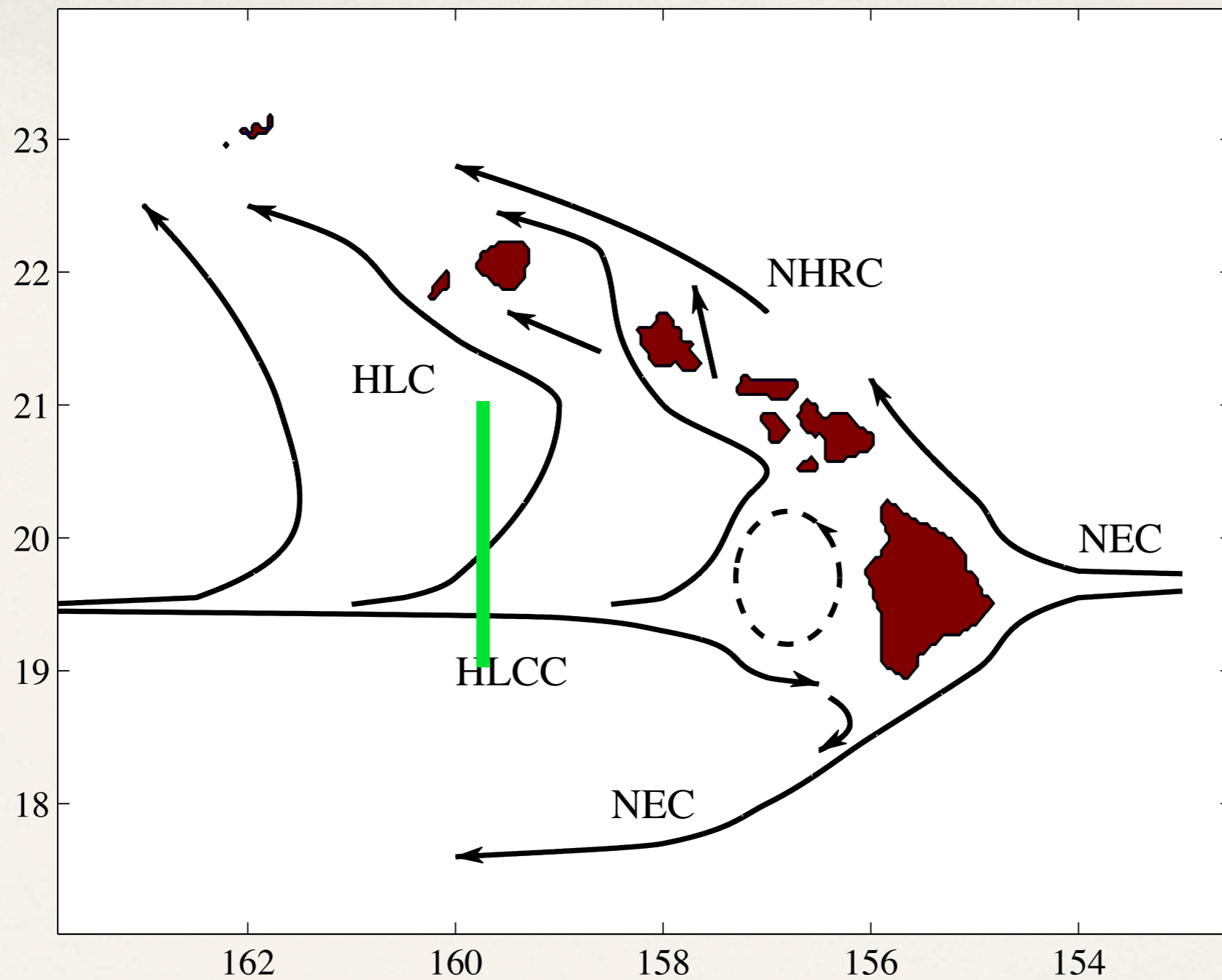
$$\mathbf{K} = (\mathbf{G}^T \mathbf{R}^{-1} \mathbf{G} + \mathbf{P}^{-1})^{-1} \mathbf{G}^T \mathbf{R}^{-1}$$

- ❖ We have some measure of the ocean, \mathcal{J}
- ❖ Between the analysis and background, we have:

$$\Delta \mathcal{J} = \mathcal{J} (\mathbf{x}_b + \mathbf{K} (\mathbf{y} - \mathbf{H}\mathbf{x}_b)) - \mathcal{J} (\mathbf{x}_b)$$

- ❖ Following Langland and Baker (2004), Errico (2007), first-order Taylor Expansion:

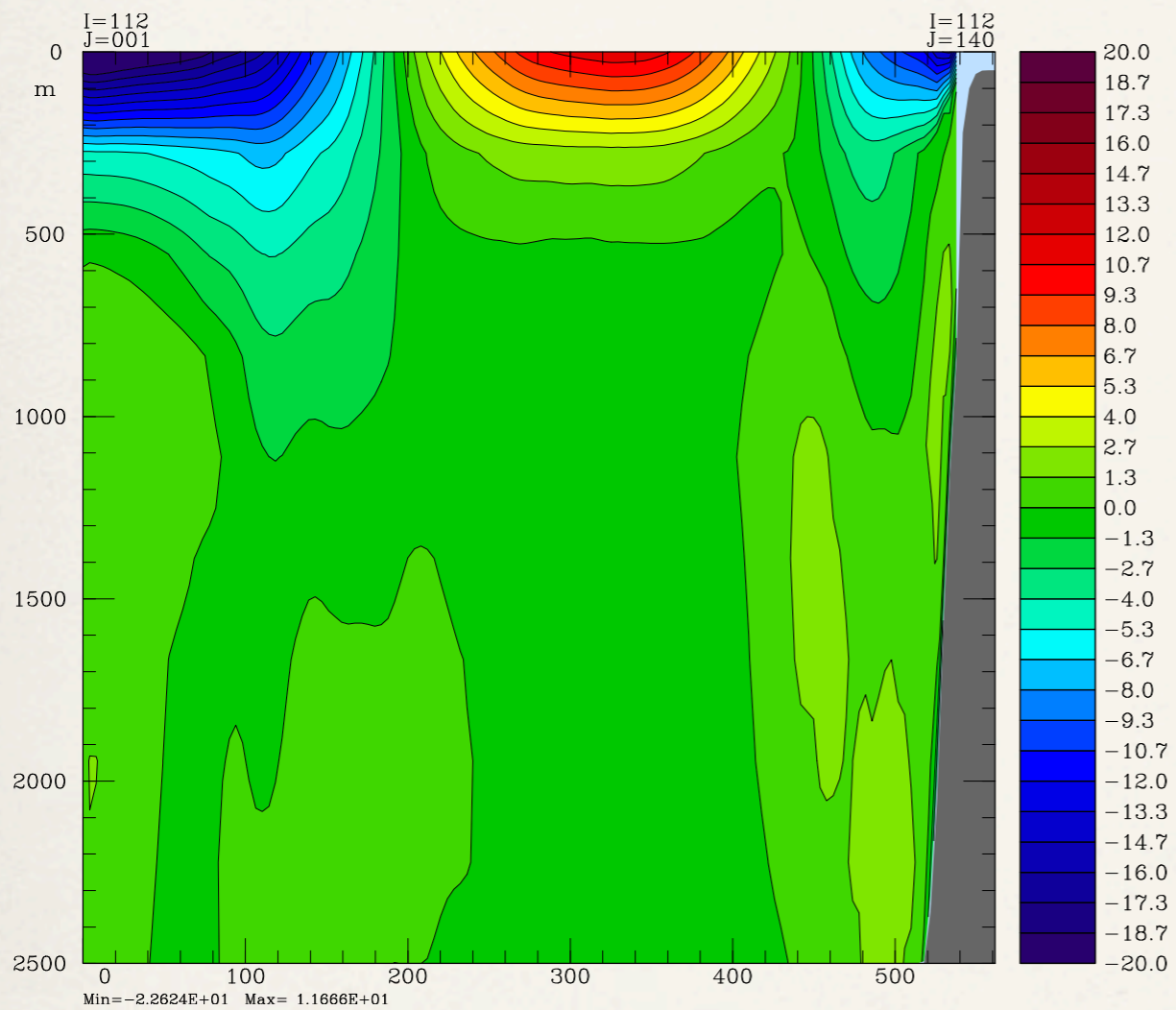
$$\Delta \mathcal{J} = 2 (\mathbf{y} - \mathbf{H}\mathbf{x}_b)^T \mathbf{K}^T \mathbf{M}_b^T \frac{\partial \mathcal{J}}{\partial \mathbf{x}_b}$$



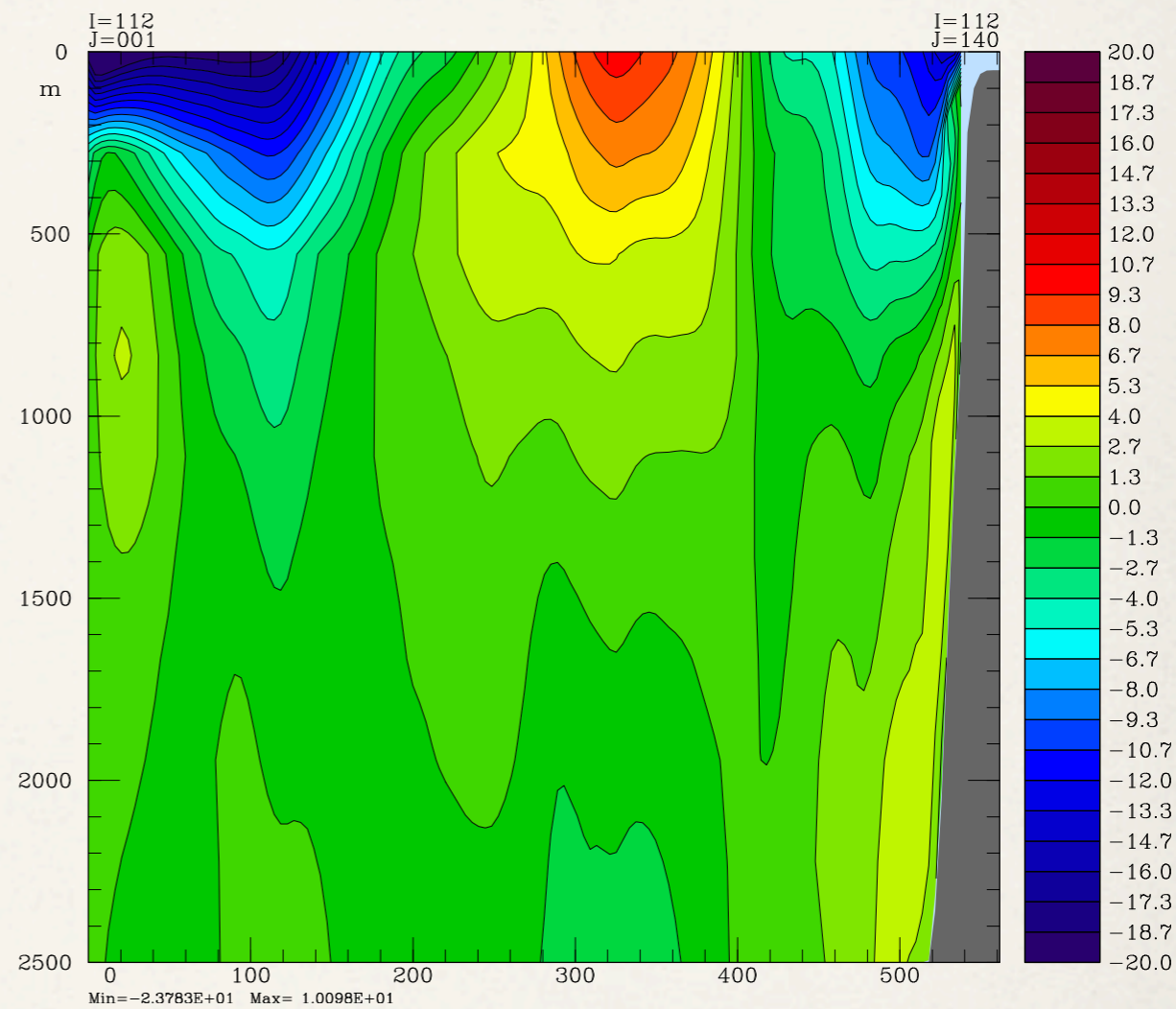
$$\mathcal{J} = \frac{1}{T \cdot 10^6} \int_T \int_S \int_{-Z}^0 u \Delta y dz ds dt$$

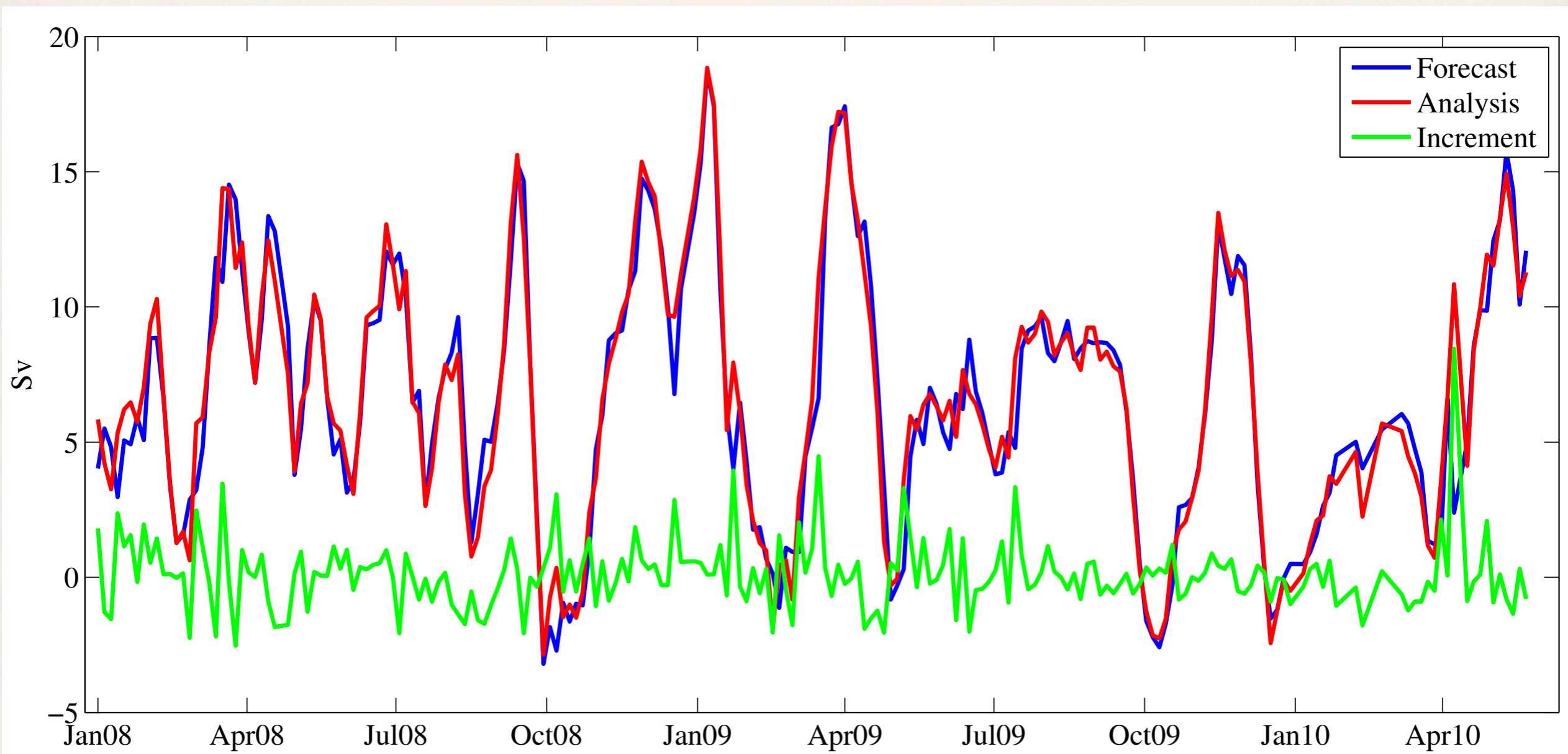


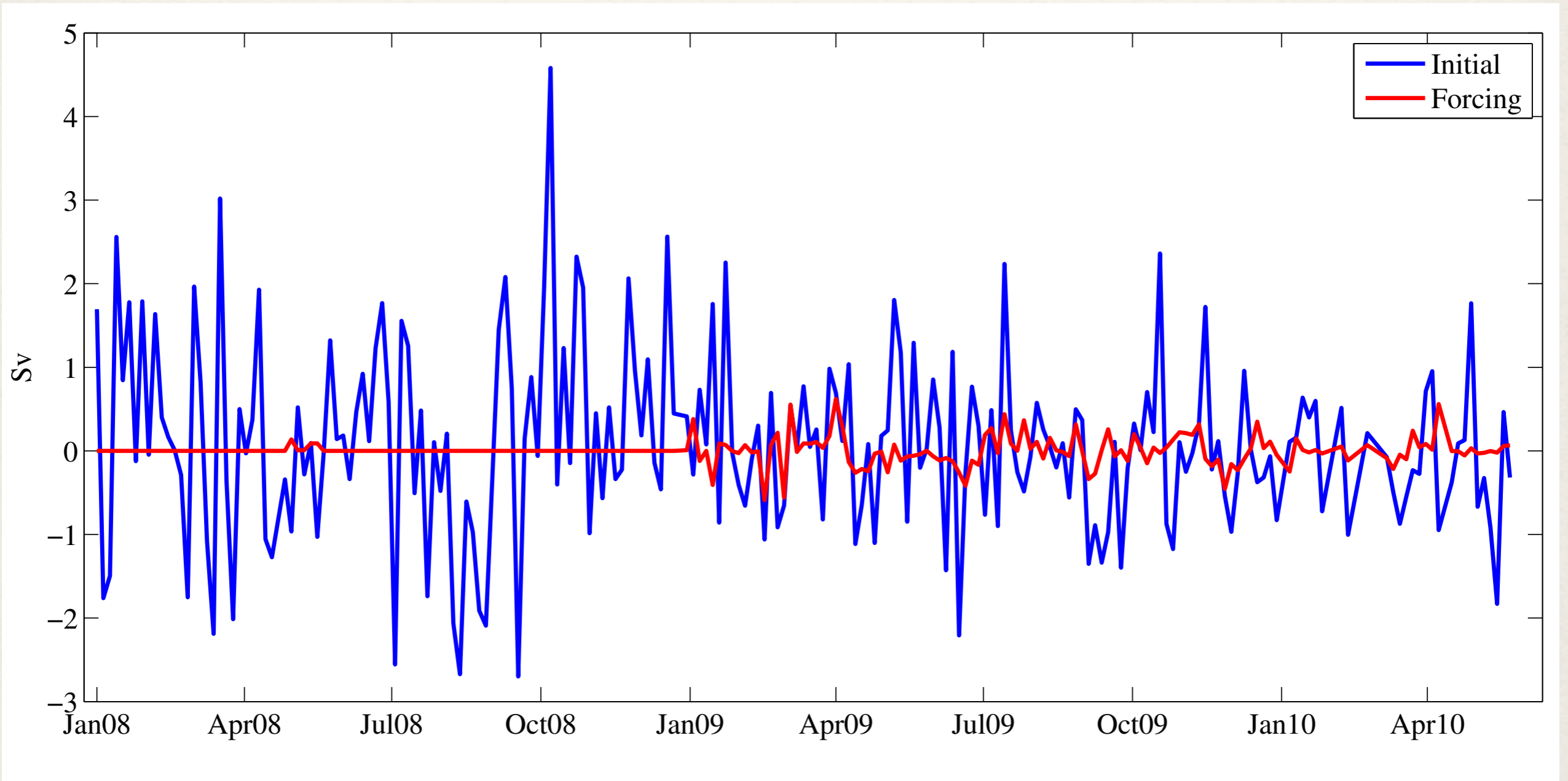
Forecast

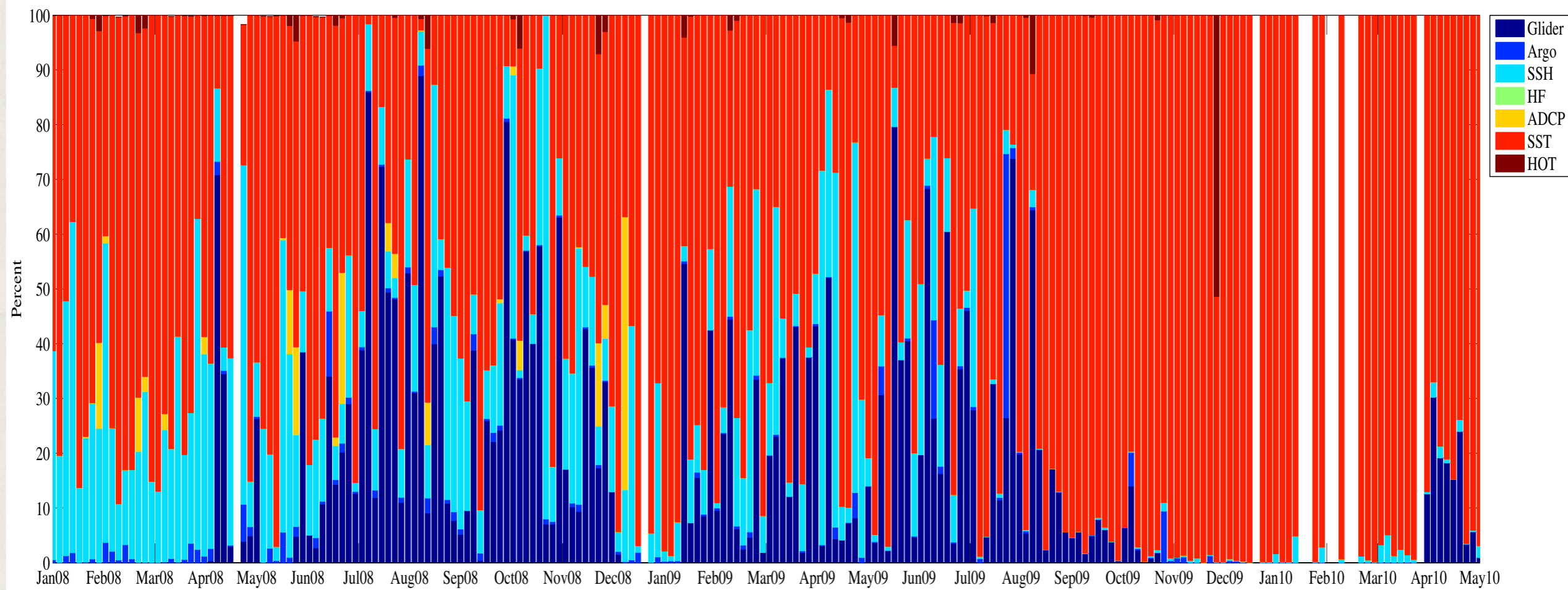
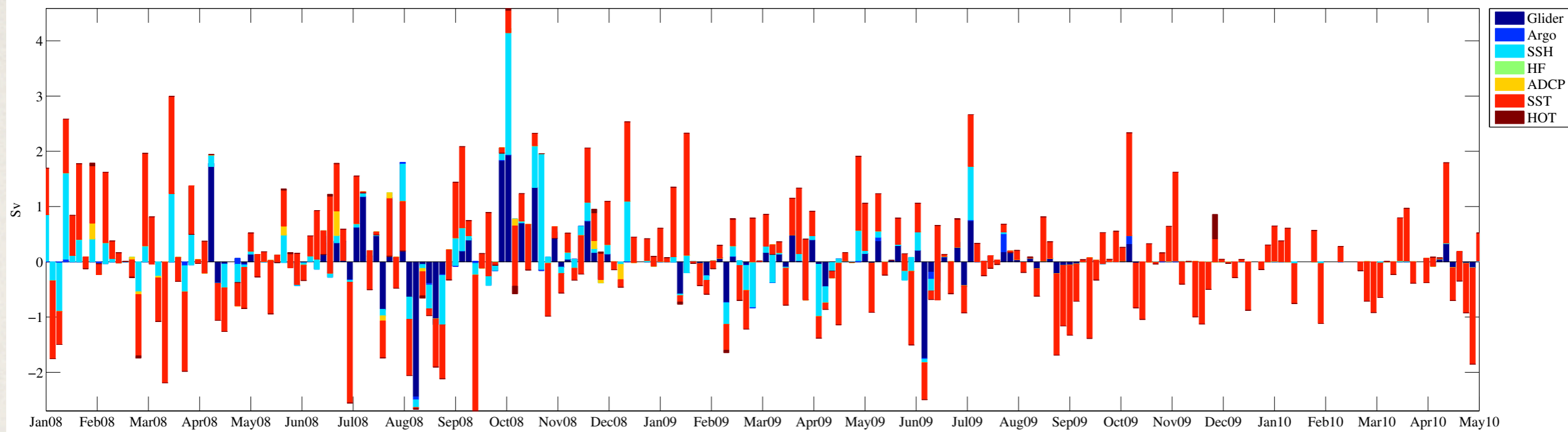
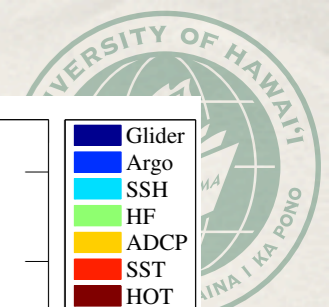


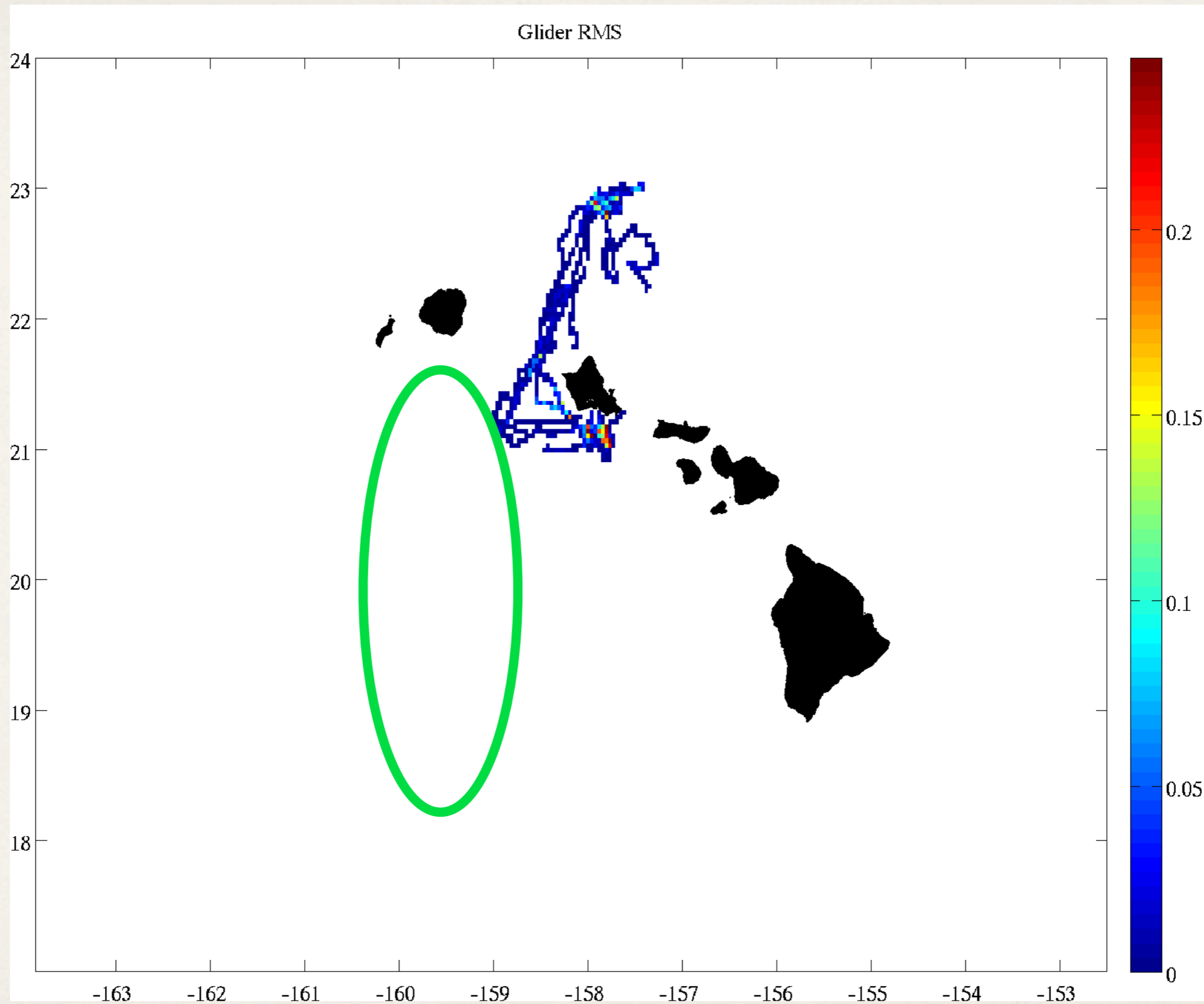
Analysis





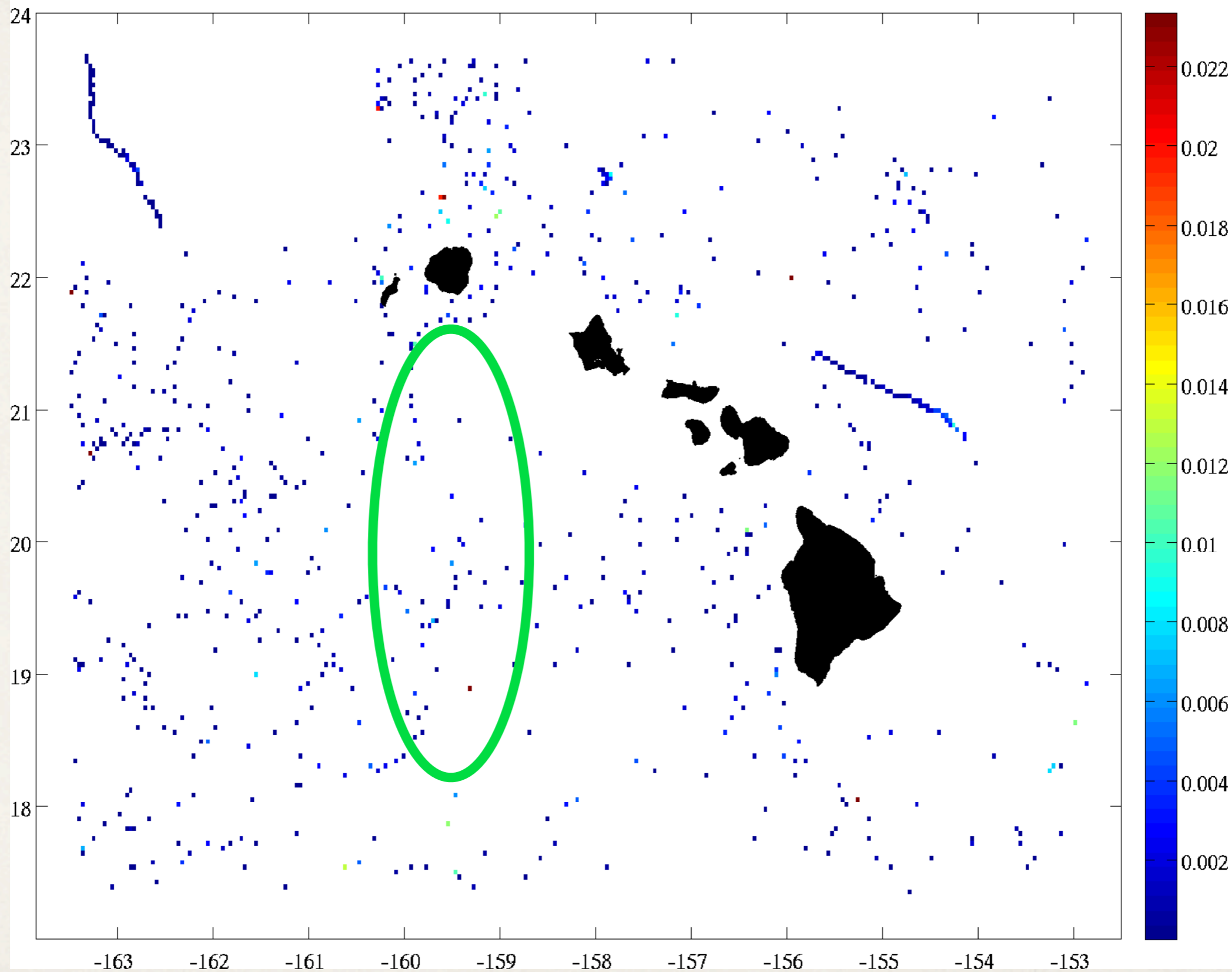


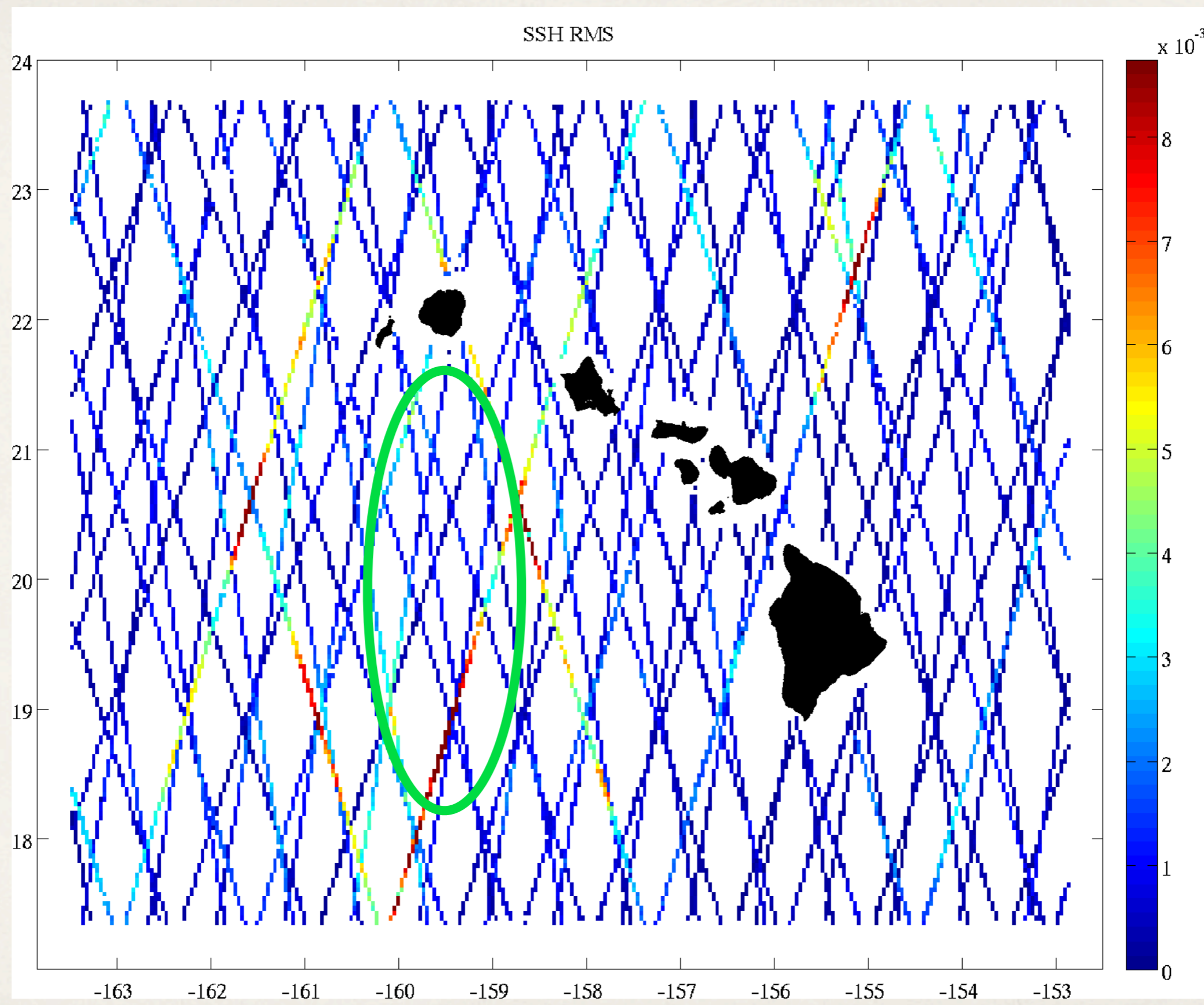


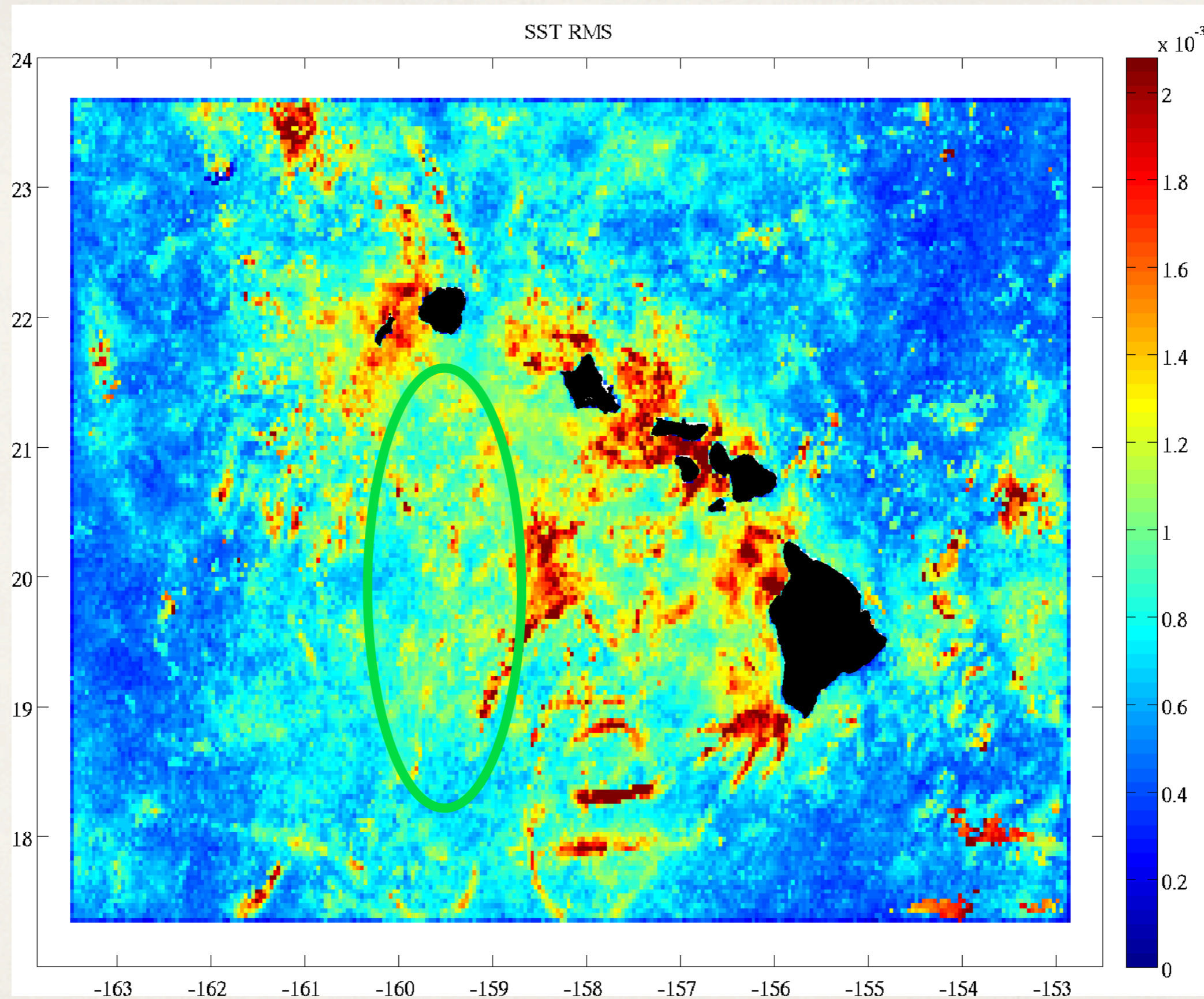




Argo RMS









	Temp	Salt	Depth
Glider	9%	5%	0-150m
	32%	42%	150-500m
Profiles	27%	7%	0-150m
	50%	15%	150-500m
		Velocity	
ADCP		21%	0-150m
		59%	150-500m



Innovation Vector: $\mathbf{d} = \mathbf{y} - \mathbf{H}\mathbf{x}_b$

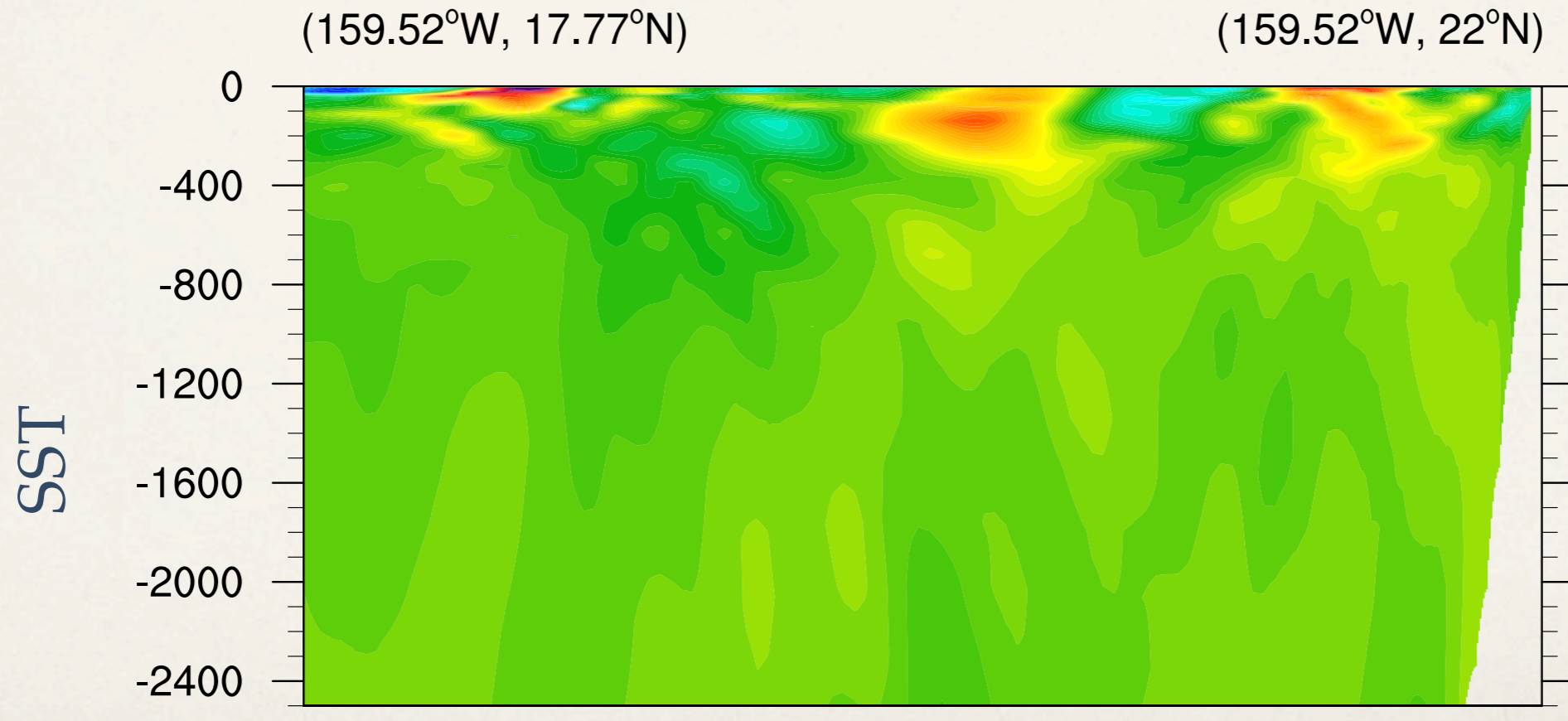
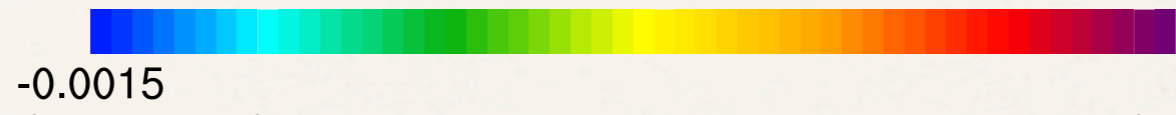
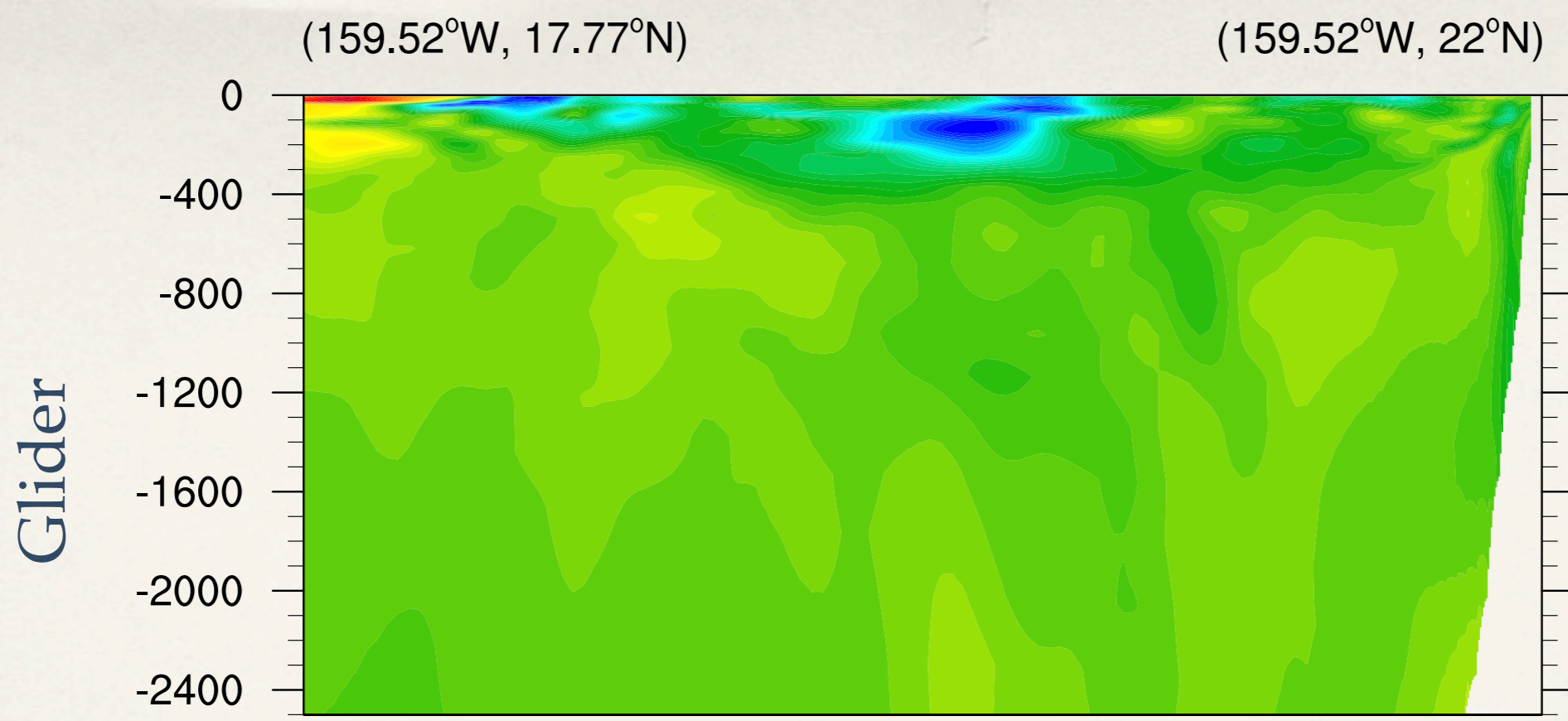
Errors Given By: $\epsilon_b = \mathbf{x}_b - \mathbf{x}_t$

$$\epsilon_y = \mathbf{y} - \mathbf{H}\mathbf{x}_t$$

$$\begin{aligned} \text{Covariance: } \langle \mathbf{d}\mathbf{d}^t \rangle &= \left\langle (\epsilon_y - \mathbf{H}\epsilon_b) (\epsilon_y - \mathbf{H}\epsilon_b)^T \right\rangle \\ &= \mathbf{R} + \mathbf{G}\mathbf{P}\mathbf{G}^T \end{aligned}$$

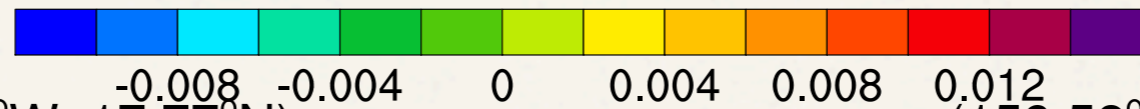
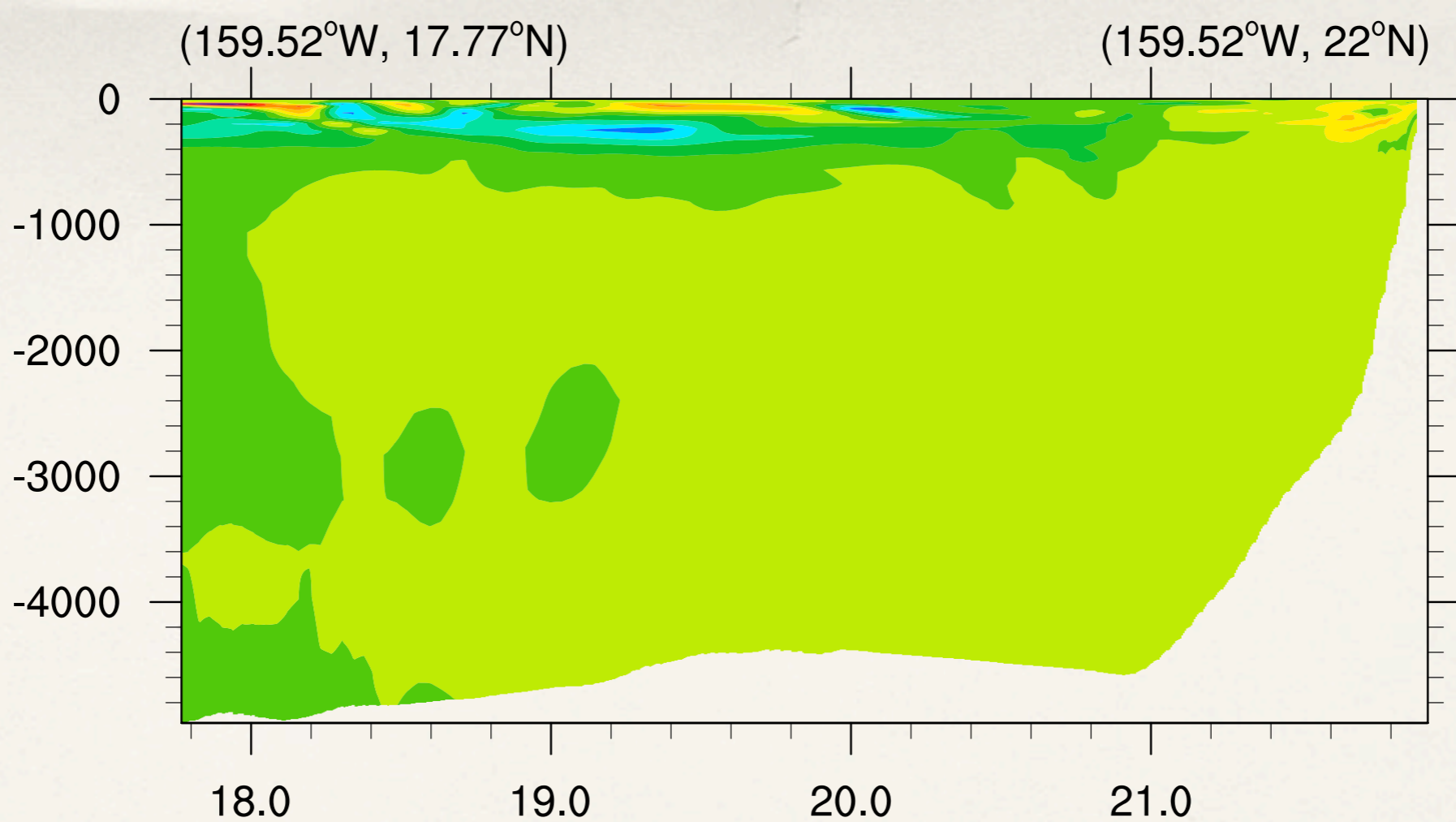
What about a single observation?

$$\mathbf{r}_m = \mathbf{M}\mathbf{P}\mathbf{G}^T \delta(y_i)$$

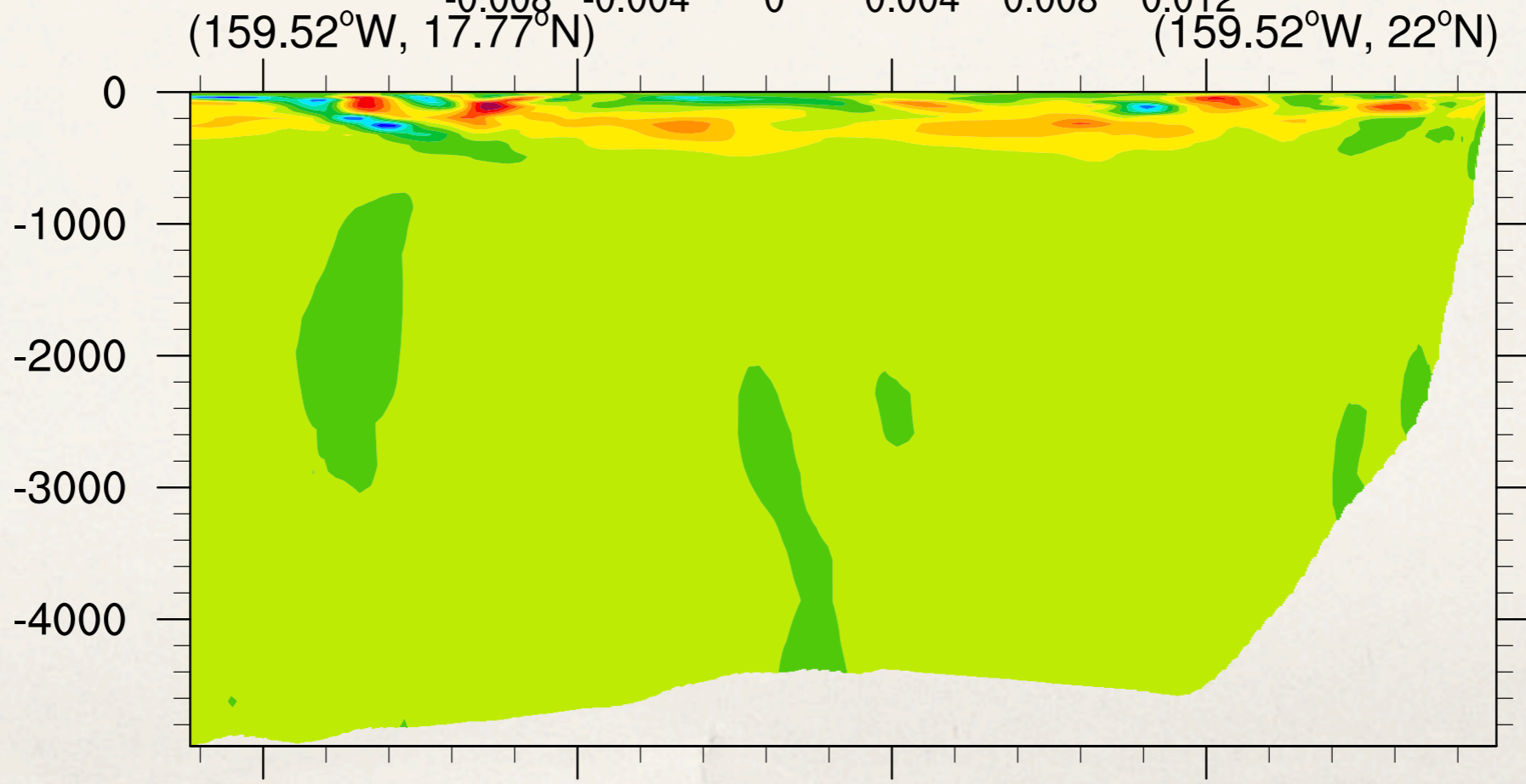


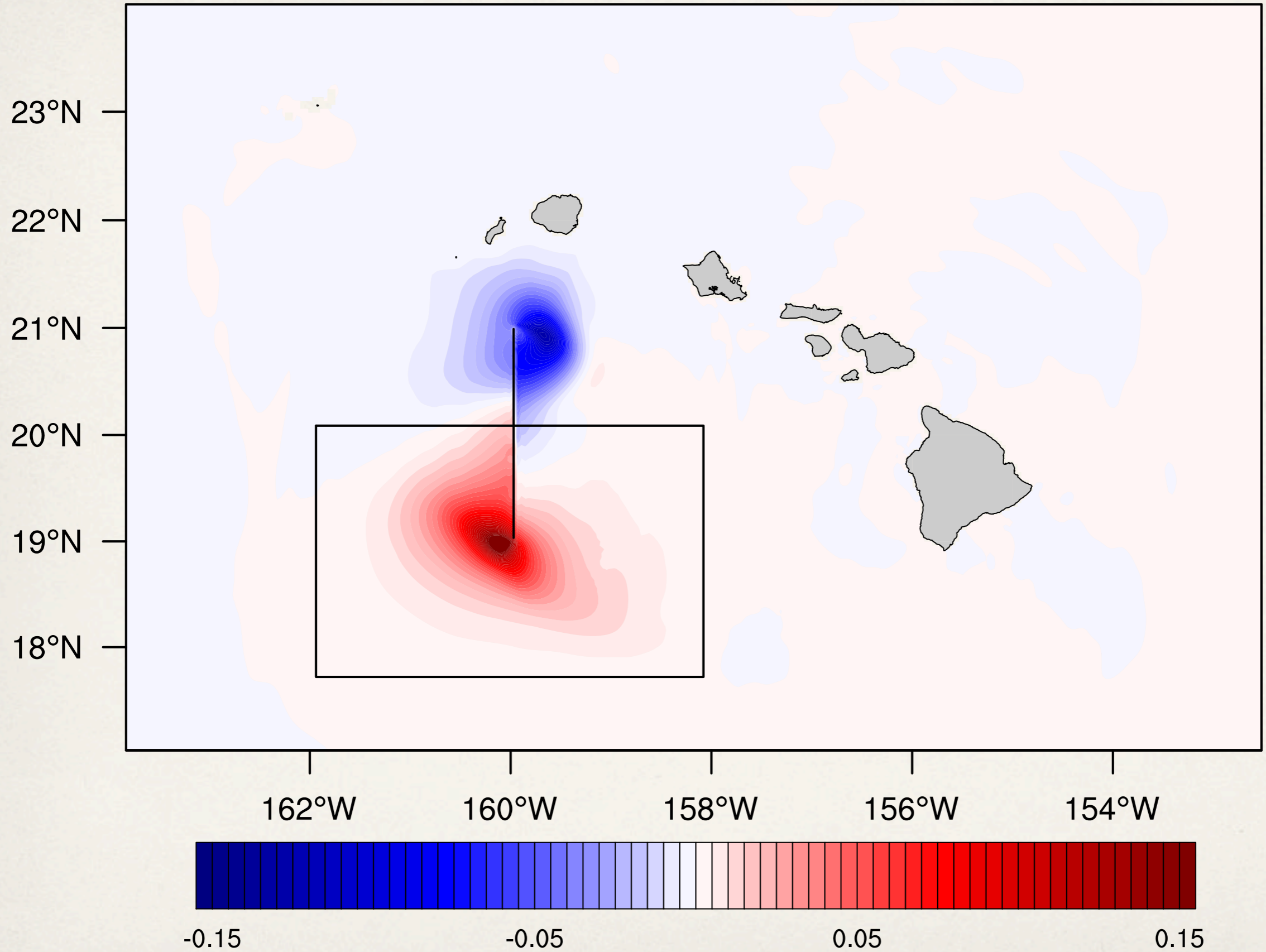


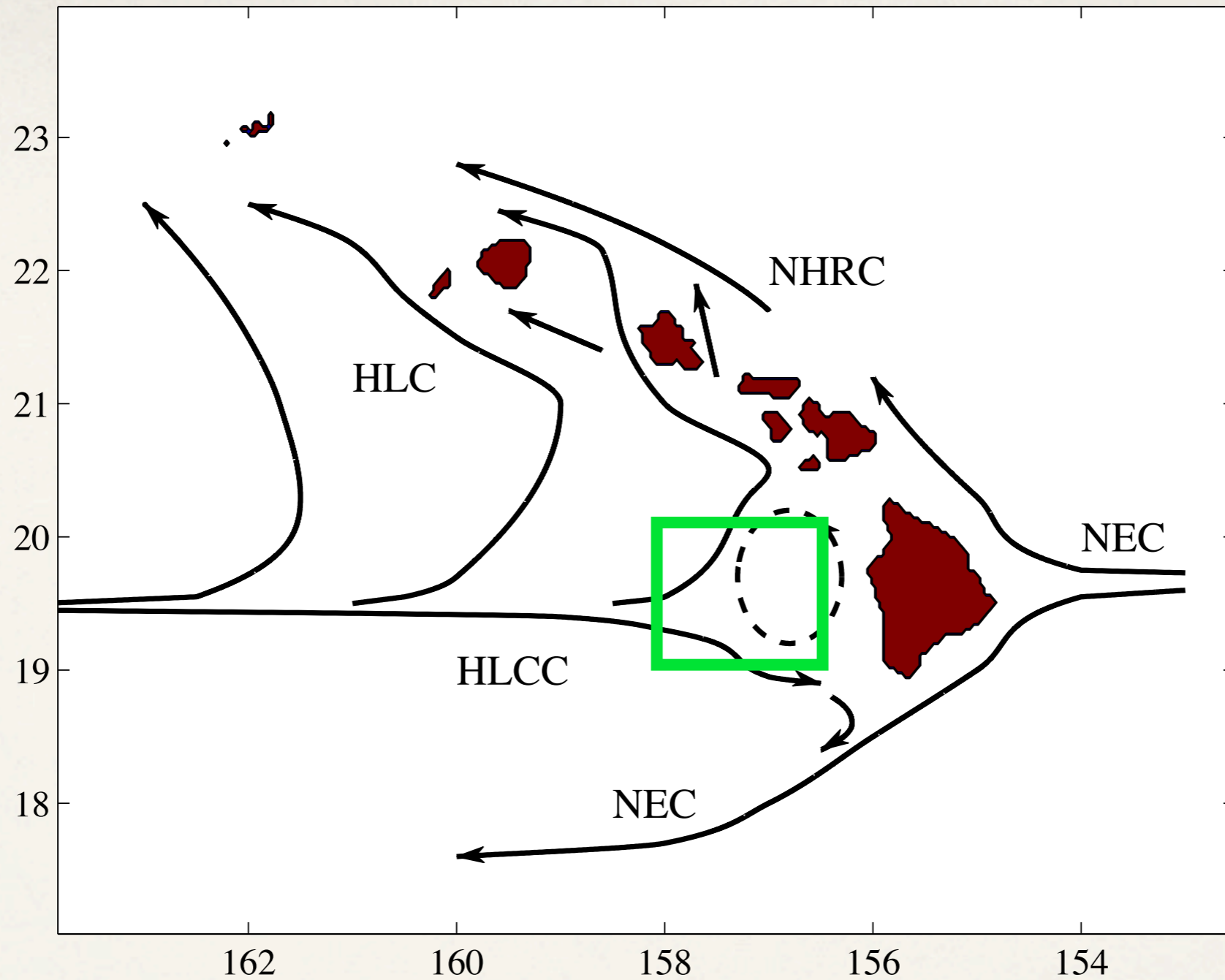
Glider



SST





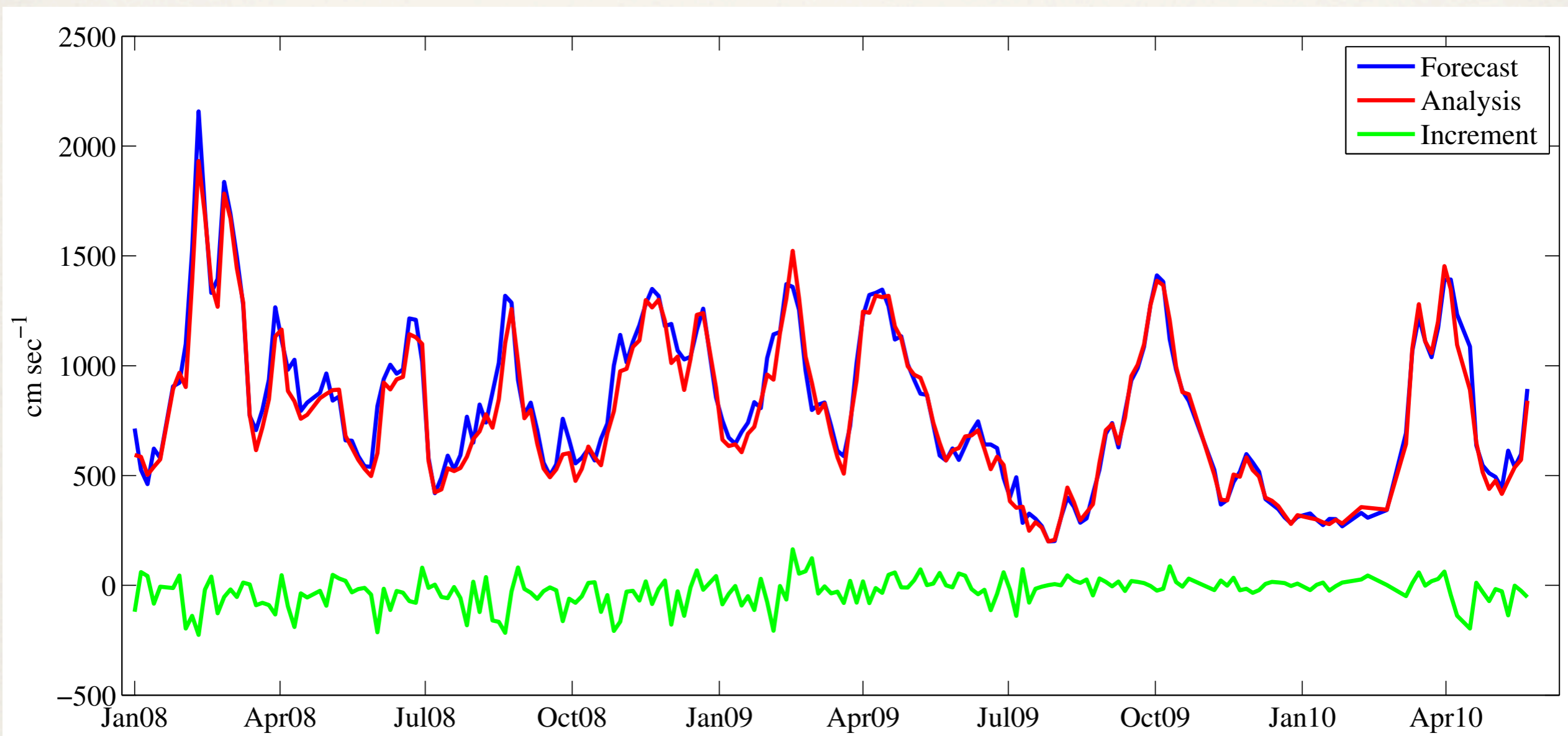


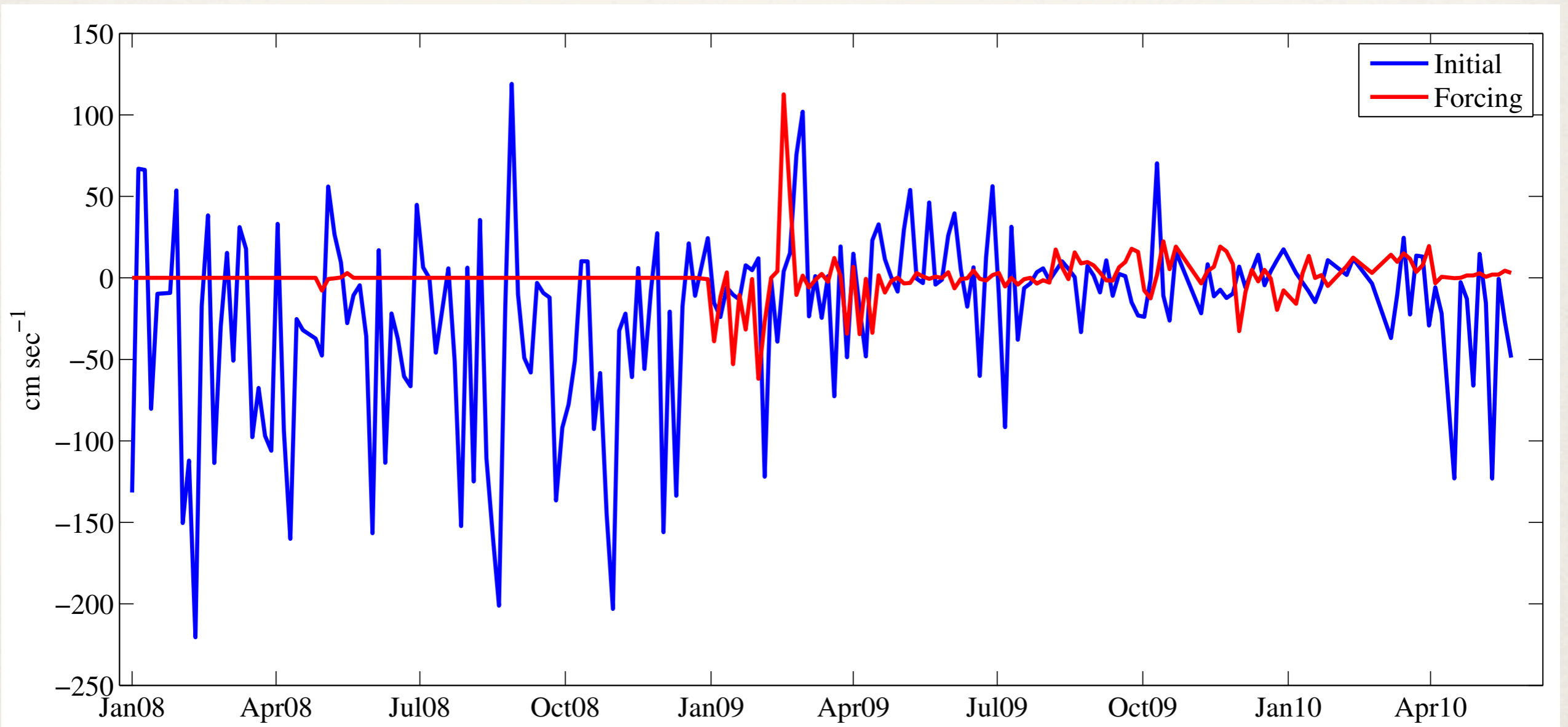
$$\mathcal{J} = \frac{1}{TZS} \int_T \int_S \int_{-Z}^0 (u^2 + v^2) dz ds dt$$

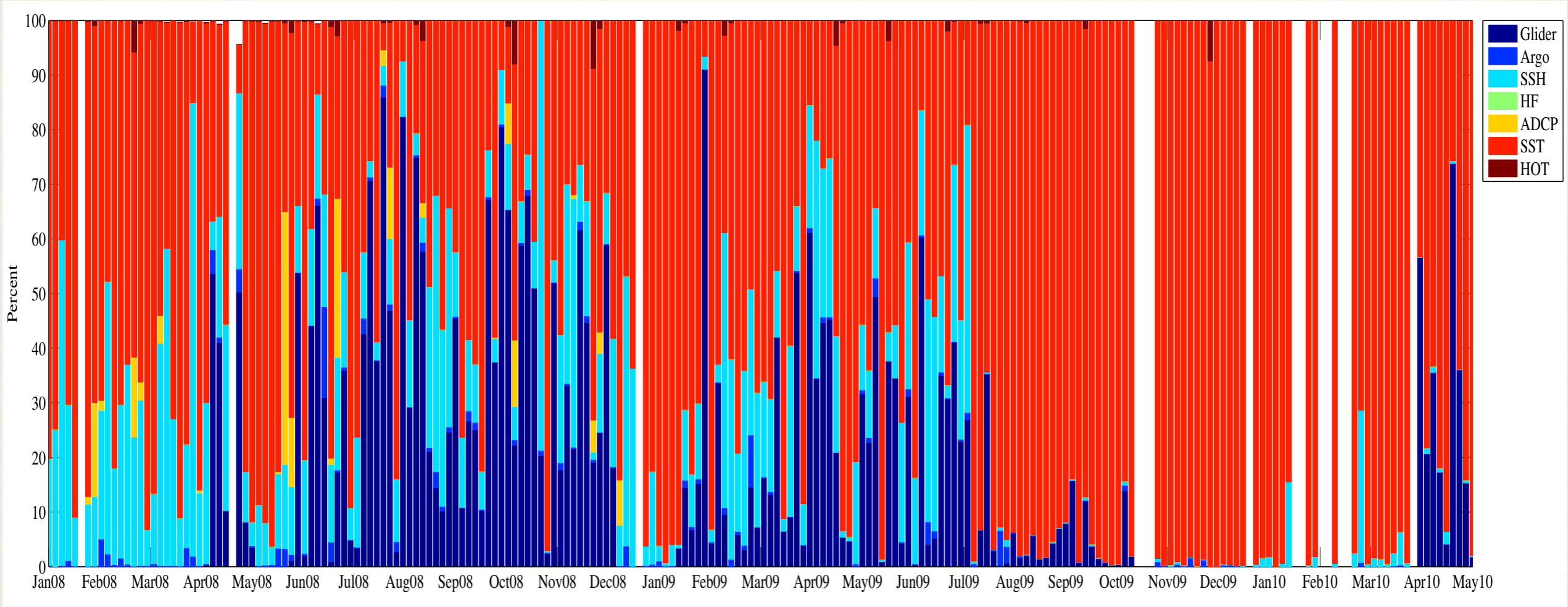
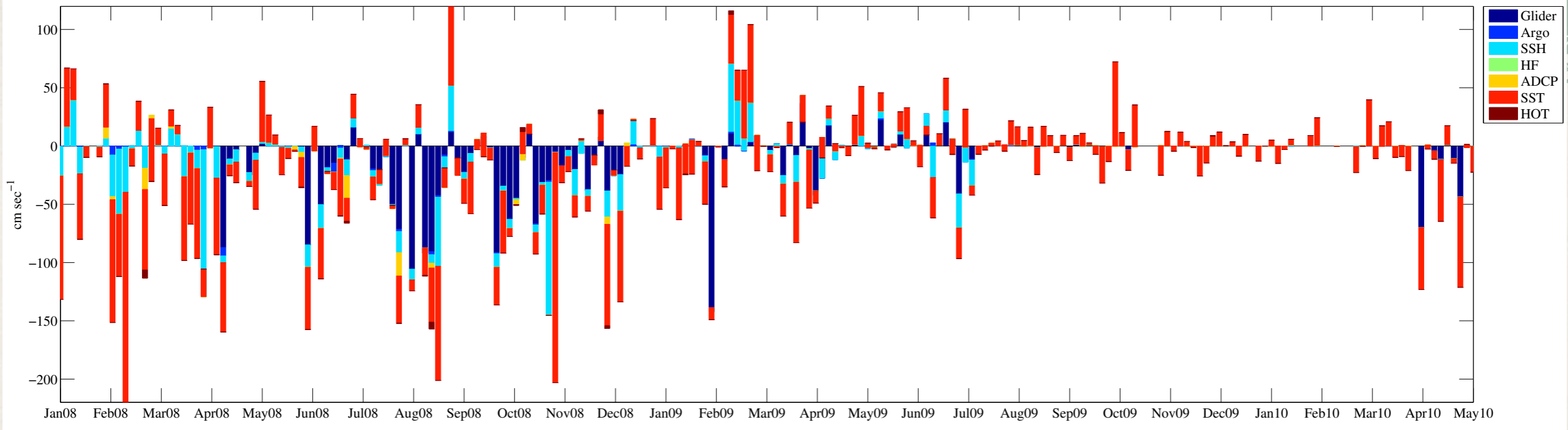
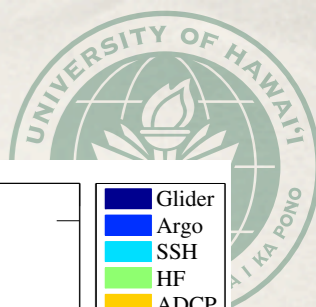


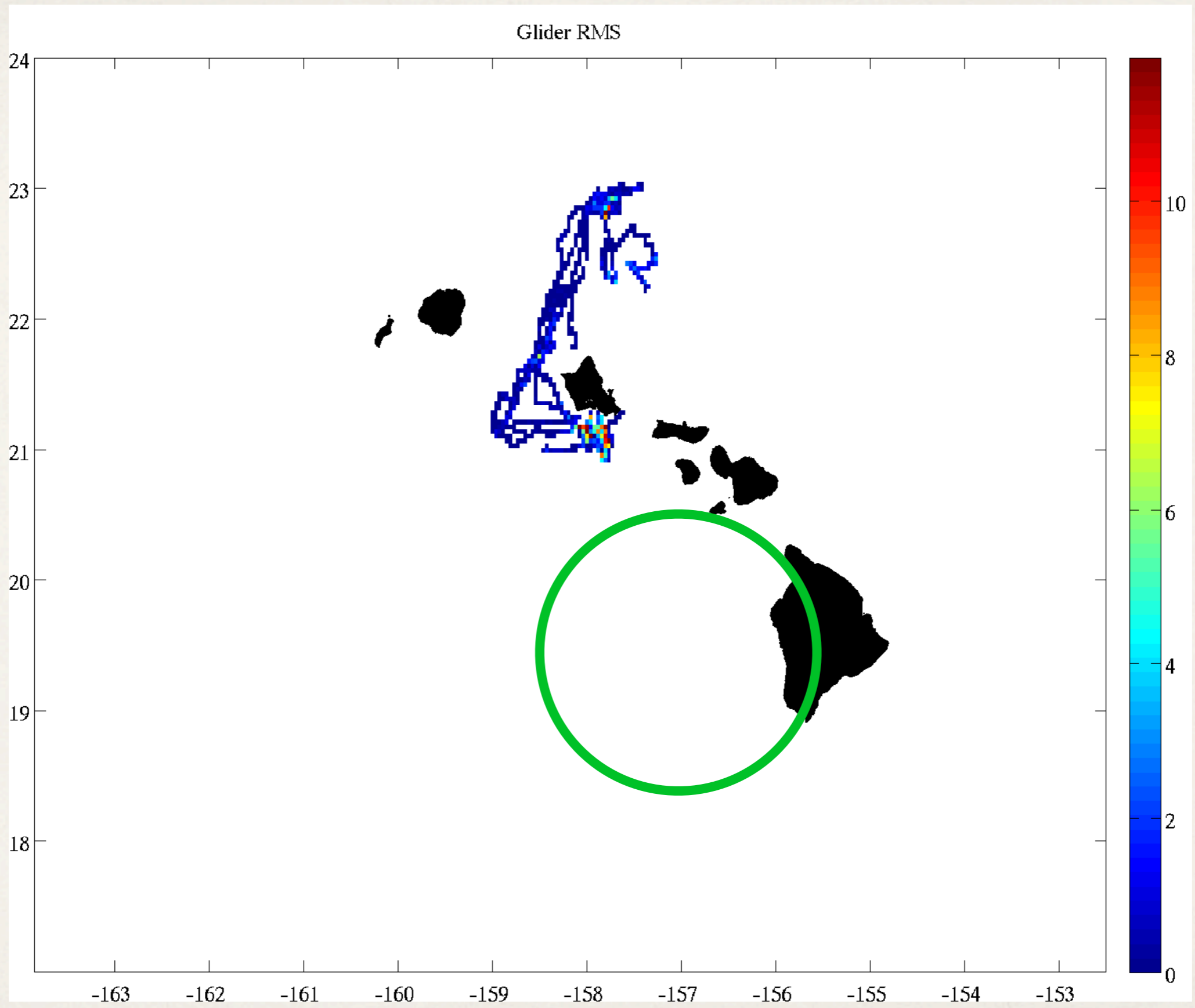
- ✦ Requires higher-order adjoint (Errico, 2007, Gelaro *et al.*, 2007)

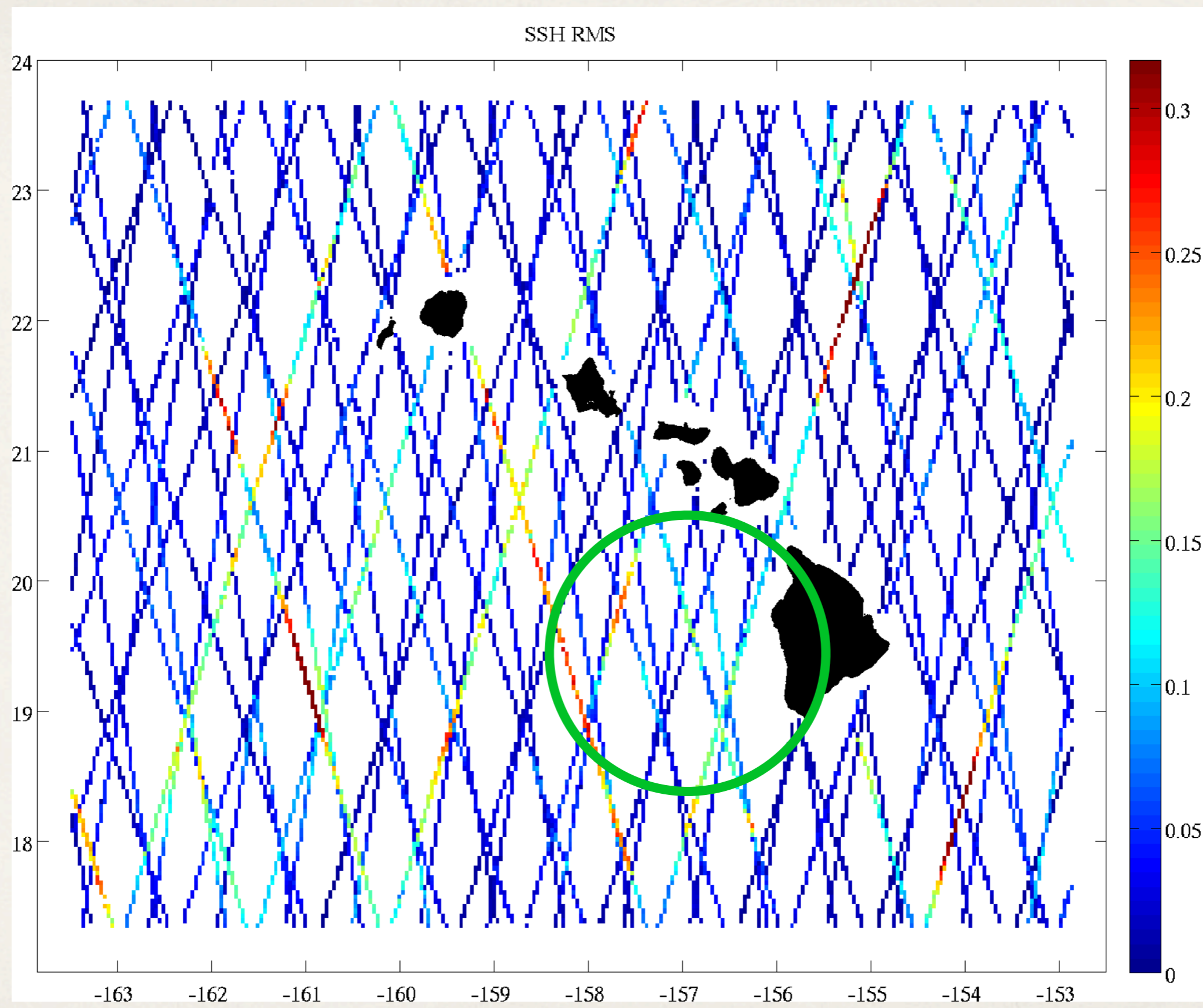
$$\Delta \mathcal{J} = (\mathbf{y} - \mathbf{H}\mathbf{x}_b)^T \mathbf{K}^T \left[\mathbf{M}_b^T \frac{\partial \mathcal{J}}{\partial \mathbf{x}_a} + \mathbf{M}_a^T \frac{\partial \mathcal{J}}{\partial \mathbf{x}_b} \right]$$

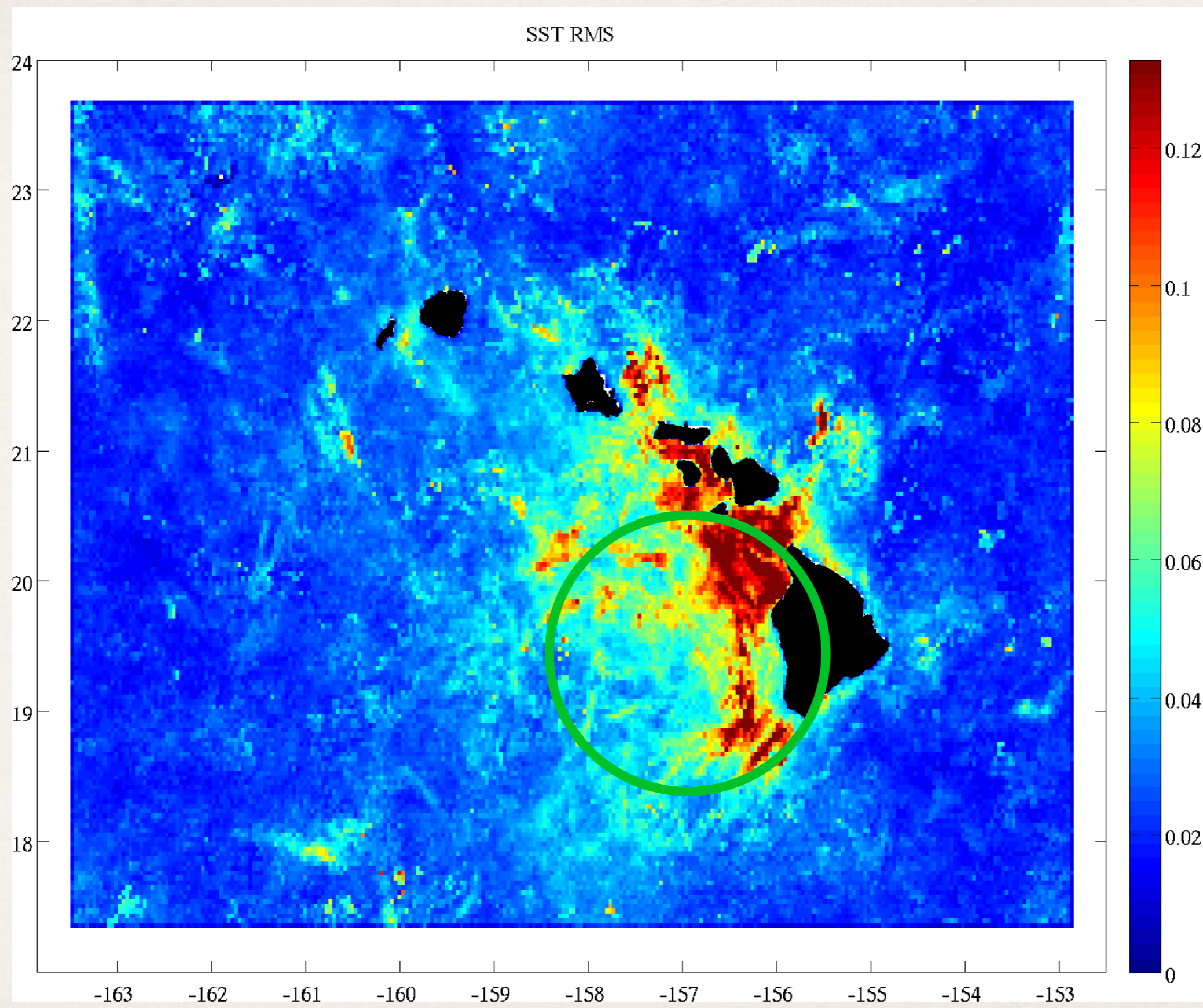




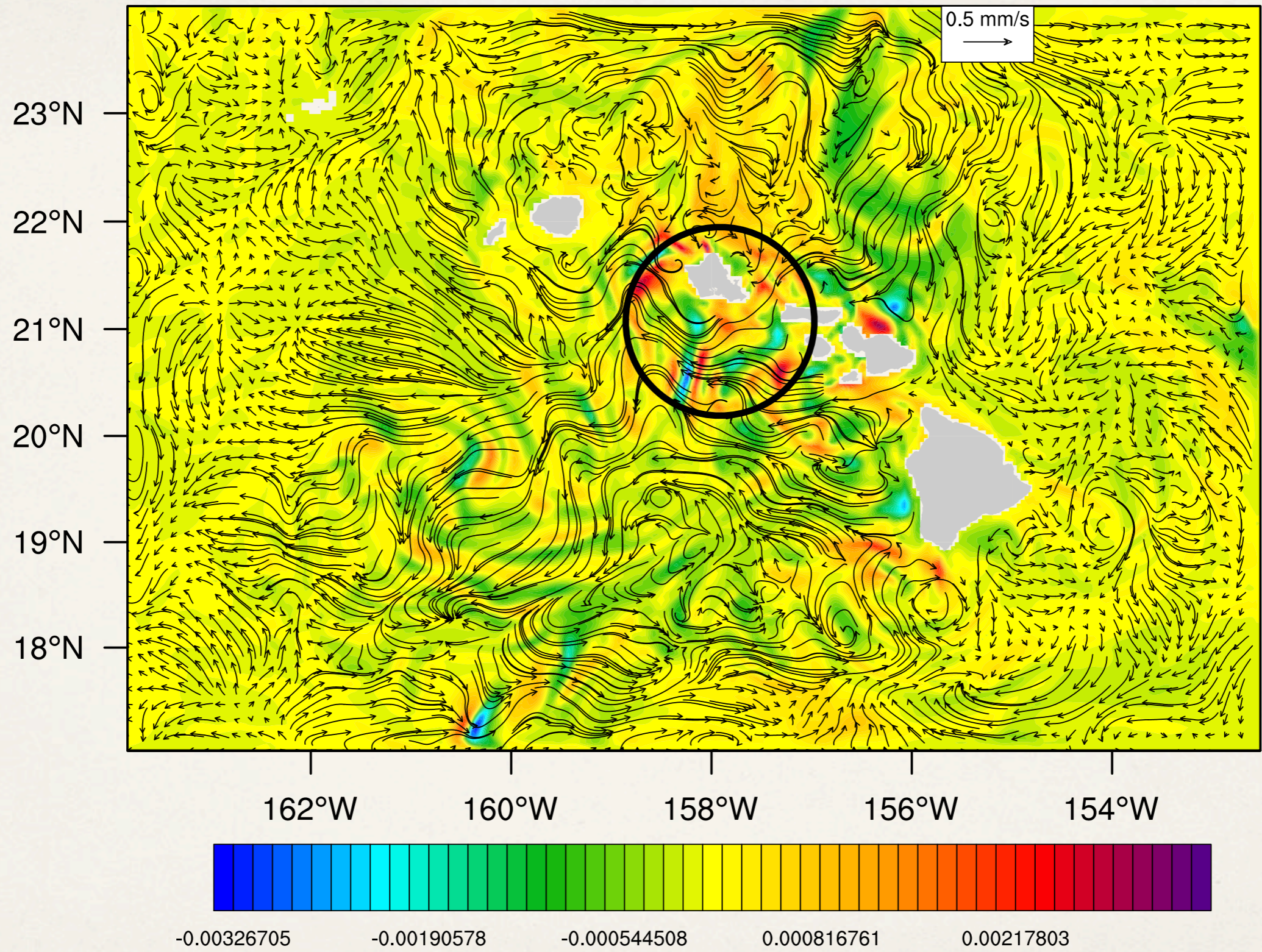




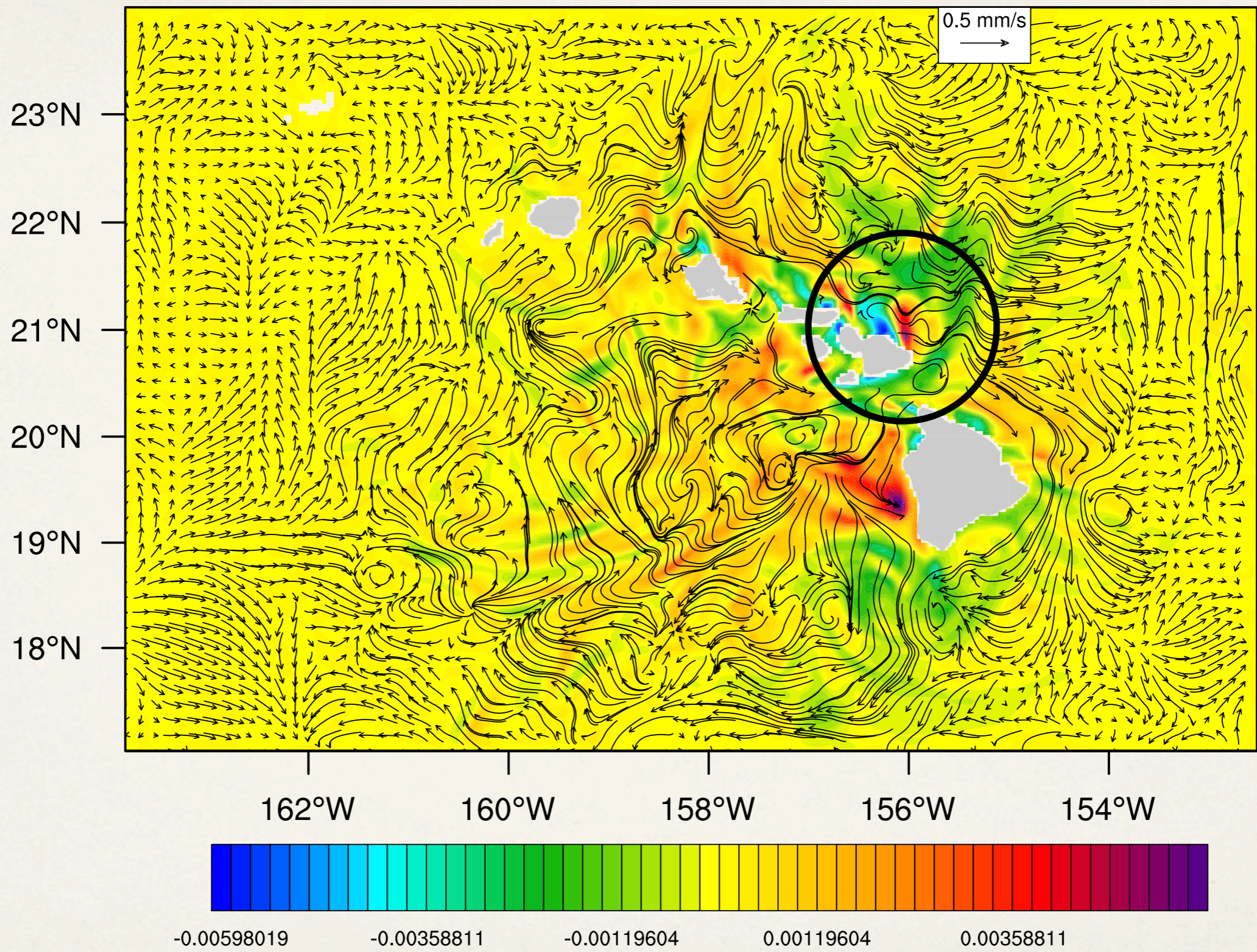


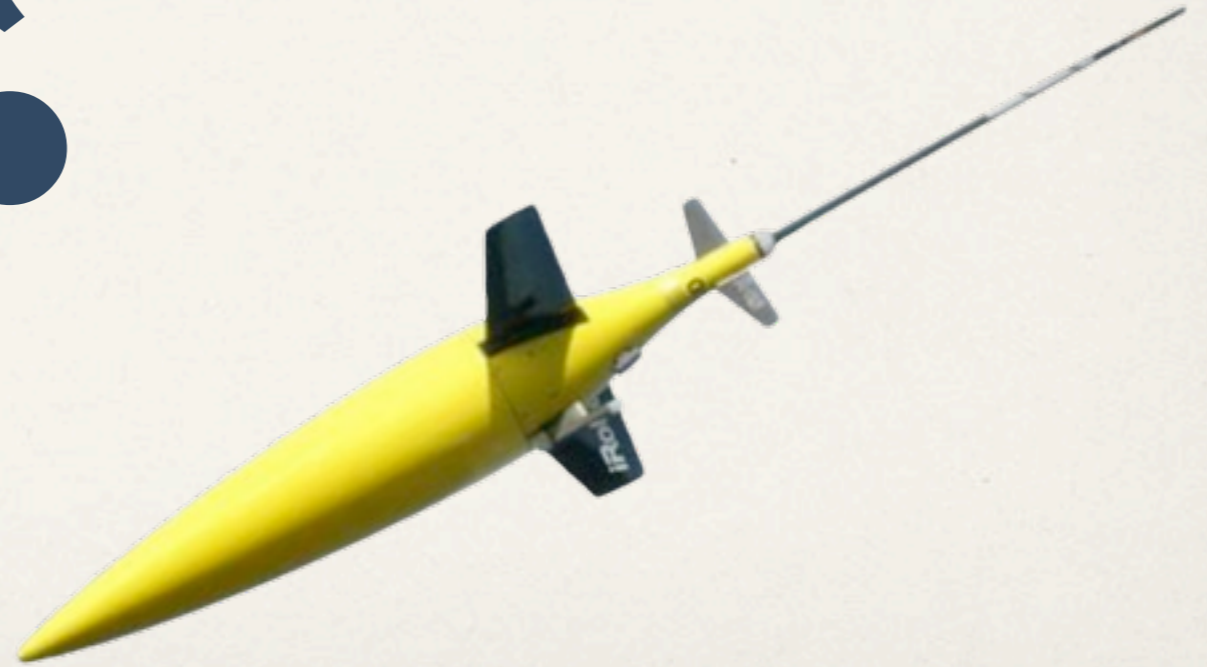


Glider



SST







HOT 1-217 stn 2

